

ADAKITES FROM ECUADOR: PRELIMINARY DATA

Michel MONZIER ⁽¹⁾, Jean-Philippe EISSEN ⁽²⁾, Jo COTTEN ⁽³⁾, Minard L. HALL ⁽⁴⁾,
Claude ROBIN ⁽¹⁾ and Pablo SAMANIEGO ⁽⁴⁾

- (1) ORSTOM, Apartado postal 17-11-6596, Quito, Ecuador (e-mail: monzier@orstom.ecx.ec)
 (2) ORSTOM, Centre de Brest, BP 70, 29280 PLOUZANE cedex, France (e-mail: eissen@ifremer.fr)
 (3) Département des Sciences de la Terre, CNRS - URA n°1278, Université de Bretagne Occidentale,
 B. P. 809, 29285 Brest, France (e-mail: Jo.Cotten@univ-brest.fr)
 (4) Instituto Geofísico, Escuela Politécnica Nacional, Apartado postal 17-01-2759, Quito, Ecuador
 (e-mail: mhall@instgeof.ecx.ec)

KEY WORDS: volcanology, geochemistry, adakites, Ecuador

INTRODUCTION

Adakites are andesites-dacites-rhyolites (but mainly dacites) with peculiar characteristics such as relatively high Al and Na contents, very low Y and Yb contents, and high La/Yb and Sr/Y ratios, that are not associated with parental basalts. They are usually considered to be derived from the partial melting of metamorphosed basalts (amphibolite and eclogite) at high pressure, leaving garnet or amphibole in the residues, thus giving them their unusual geochemical characteristics (Defant and Drummond, 1990; Defant, et al., 1991; Drummond and Defant, 1990; Atherton and Petford, 1993; Defant and Drummond, 1993; Sajona et al., 1993). It is generally accepted that adakitic magmas form where young (<25 Ma), hot oceanic crust is subducted and melts at 23-26 kbar (75-85 km) and 700-775°C (Drummond and Defant, 1990), i.e. near the corner of the mantle wedge. Accordingly, adakitic volcanoes should lie quite close (\approx 80 km) to the trench. However, Atherton and Petford (1993) presented evidence that the partial melting of basaltic material of the lower crust at \approx 50 km depths by newly underplated hot magmas (generated in the underlying asthenospheric wedge) should be considered as an additional way to generate adakite, when 1/ subducting oceanic crust is too old and cold, 2/ volcanism is very far from the trench, and 3/ local heat flow is high.

ECUADORIAN ADAKITES

Young oceanic lithosphere (<20 Ma) formed along the Cocos-Nazca spreading axis is being subducted under central and northern Ecuador since at least the early Pliocene with a rate of convergence of \approx 8 cm/y (Fig. 1). The Benioff plane dips \approx 25° eastward, attaining \approx 130-150 and \approx 150-175 km depths under the Quaternary volcanoes of the Ecuadorian Western and Eastern Cordilleras, respectively (Winter, 1990). The Eastern Cordillera constitutes the eastern edge of the Andean Block (AB), a wedge between the Nazca (NAZ) and South American (SAM) plates presently affected by an E-W compressive state of stress (Ego et al., 1995). The collision-subduction of the Carnegie Ridge in front of Ecuador (Pennington, 1981) or, more certainly, the oblique NAZ-SAM convergence in the Ecuador-Colombia area (Ego et al., 1995) would explain the northeastward motion of the AB, along the great NE-SW right-lateral transpressive faults of the East Andean Front Fault Zone.

ORSTOM (French Scientific Research Institute for the Development in Cooperation) and the Geophysical Institute of the National Polytechnical School of Quito are collaborating since 1994 in a joint volcanological program which includes studies on several volcanoes of Ecuador, namely Cotacachi, Fuya Fuya, Mojanda, Pululahua, Cayambe, Cotopaxi, Tungurahua, Sangay... (Fig. 2) We present here the first results in geochemistry (see also Robin et al. and Samaniego et al., this volume). To date, all the rocks sampled within the framework of this joint program have SiO₂ contents > 55%; they are mainly medium-K basic andesites to rhyolites, and a few high-K acid andesites to rhyolites (Fig. 3; all values LOI free and recalculated to 100%). For comparison, basaltic rocks clearly predominate in a typical

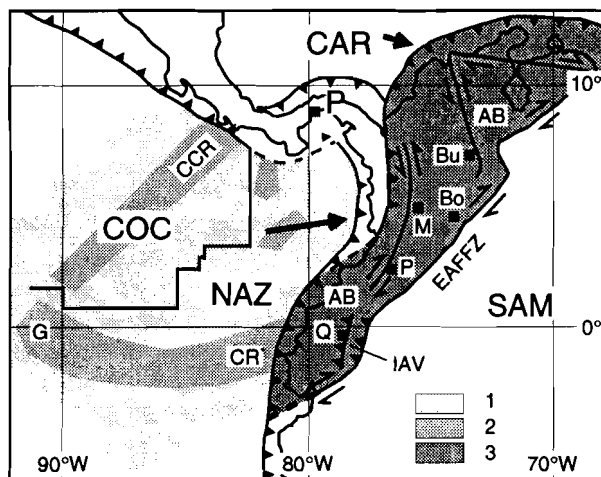
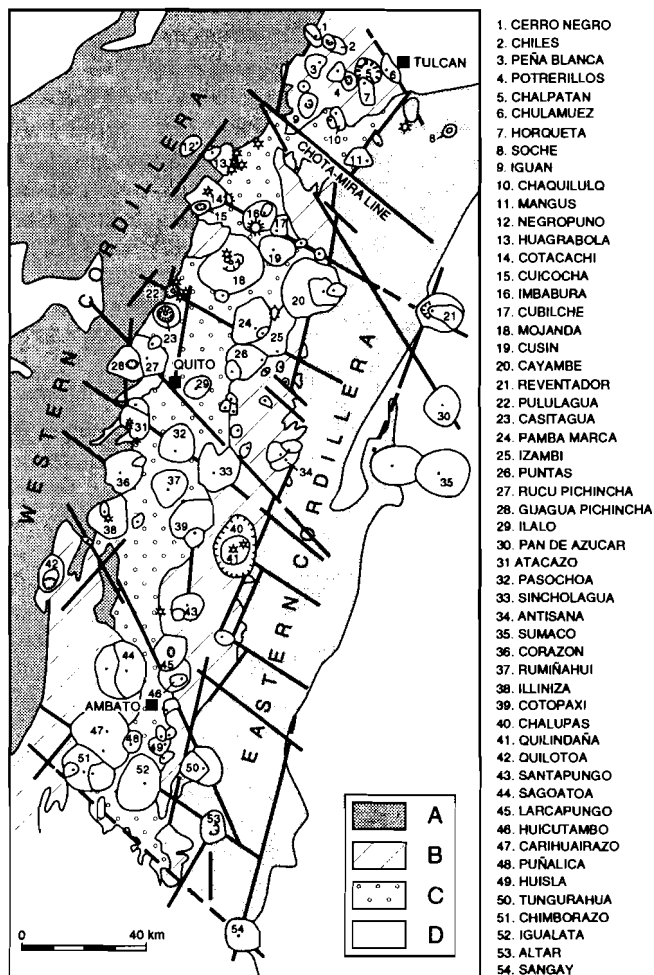


Fig. 1

Geodynamics of the Andean Block.
 COC= Cocos Plate;
 NAZ= Nazca Plate;
 CAR= Caribbean Plate;
 1= Galapagos Gore; 2= oceanic ridges:
 CCR= Cocos Ridge, CR= Carnegie
 Ridge; 3= Andean Block (AB);
 G= Galapagos Islands.
 Large arrows= NAZ and CAR plate
 convergence with respect to SAM fixed.
 IAV= Inter Andean Valley;
 EAFZZ= East Andean Front Fault Zone;
 Q= Quito, P= Popayan, M= Manizales,
 Bo= Bogota, Bu= Bucaramanga, P=
 Panama. After Pennington (1981) and
 Ego et al. (1993).

Fig. 2

Quaternary volcanoes
 and main tectonic lineations
 of the Ecuadorian Andes
 from 1° N to 2° S
 (from Barberi et al., 1988).
 A= Western Cordillera: Macuchi
 Formation and Cretaceous-Eocene
 clastic deposits;
 B= undifferentiated basal
 volcanic complex (Late Miocene-
 Quaternary);
 C= volcanic and volcano-sedimen-
 tary deposits filling the Interandean
 Depression;
 D=Eastern Cordillera: mainly
 metamorphites.



island-arc, such as the New Hebrides arc (Southwest Pacific; Monzier et al., in prep.). A marked increase of K₂O content is also obvious in the volcanic products from the Western Cordillera to the Eastern Cordillera, an observation already noted by Barberi et al. (1988). Another striking difference between the rocks from the two cordilleras is the value of the Ba/La ratio (LILE / HFSE), >49 in the Western Cordillera and <49 in the Eastern Cordillera (Cayambe, Chacana, Cotopaxi, Tungurahua), an observation which may reflect a greater slab contribution in the magmas of the Western Cordillera. Diagrams of Sr/Y vs. Y and La/Yb vs. K₂O (Fig. 3) emphasize the very high Sr/Y and La/Yb ratios of recent Cayambe dacites, a characteristic signature of dacitic melts derived from metamorphosed basalts, i. e. adakites. These adakites from Cayambe clearly follow a broad melting trend on the La/Yb vs. K₂O diagram, fractionation being only of minor importance. Contrarily, in the Tungurahua suite fractionation is the dominant process. Sr/Y ratios and Y contents of adakites from Cayambe and Mount St Helens (Defant and Drummond, 1993) are very similar. Lastly, it should be stressed that samples from the old part of Cayambe are not adakitic. All samples from Cushnirumi and Fuya Fuya also have a strong adakitic character, whereas only some of the rocks from Cotacachi, Pululahua, Quilotoa, Cotopaxi and Tungurahua volcanoes are adakites. Conversely, all Mojanda samples are calc-alkaline.

PRELIMINARY CONCLUSIONS

At this point of the program, it appears that adakites are common in Ecuador, in both cordilleras. They can appear either during a given period of the history of a single volcano (Cayambe: Samaniego et al., this volume) or simultaneously with calc-alkaline rocks at two close volcanic centers (Fuya Fuya and Mojanda: Robin et al., this volume), which entail some interesting volcanological questions. The question of the basaltic source, i. e. slab or lower crust melting, remains unsolved for the moment, but as ecuadorian adakitic volcanoes (especially Cayambe) lie far from the trench, it is perhaps more reasonable to contemplate a crustal origin than a slab one, even if the subducting crust is young (<20 Ma). Regional crustal modeling and heat flow measurements would be very convenient to solve this question as well as detailed studies on the mode of subduction below Ecuador and the role of the Carnegie Ridge in this process.

REFERENCES

- Atherton, M. P. and Petford, N., 1993. Generation of sodium-rich magmas from newly underplated basaltic crust. *Nature*, 362, 144-146.
- Barberi, F., Coltelli, M., Ferrara, G., Innocenti, F., Navarro, J.M. and Santacroce, R., 1988. Plio-Quaternary volcanism in Ecuador. *Geol. Mag.*, 125, 1, 1-14.
- Ego, F., Sebrier, M., Lavenu, A., Yopez, H. and Eguez, A., 1993. Quaternary state of stress in the northern Andes and the restraining bend model for the ecuadorian Andes. Second ISAG, Oxford (UK), 21-23/9/1993, extended abstracts, Editions de l'ORSTOM, Paris 1993, p. 89-92.
- Ego, F., Sébrier, M. and Yepes, H., 1995. Is the Cauca-Patia and Romeral fault system left or rightlateral? *Geophys. Res. Lett.*, 22, 1, 33-36.
- Defant M. J. and Drummond, M. S., 1990. Derivation of some modern arc magmas by melting of young subducted lithosphere. *Nature*, 347, 662-665.
- Defant M. J. and Drummond, M. S., 1993. Mount St. Helens: potential example of the partial melting of the subducted lithosphere in a volcanic arc. *Geology*, 21, 547-550.
- Defant, M. J., Richerson, P. M., De Boer, J. Z., Stewart, R. H., Maury, R. C., Bellon, H., Drummond, M. S., Feigenson, M. D. and Jackson, T. E., 1991. Dacite genesis via both slab melting and differentiation: petrogenesis of La Yeguada volcanic complex, Panama. *J. Petrol.*, 32, 6, 1101-1142.
- Drummond, M. S. and Defant, M. J., 1990. A model for trondhjemite-tonalite-dacite genesis and crustal growth via slab melting: Archean to modern comparisons. *J. Geophys. Res.*, 95, B13, 21503-21521.
- Monzier, M., Robin, C., Eissen, J. P. and Cotten, J. (in prep.). Geochemistry of the volcanic New Hebrides Central Chain (Southwest Pacific). *J. Volcanol. Geotherm. Res.*,....
- Pennington, W. D., 1981. Subduction of the eastern Panama Basin and seismotectonics of northwestern South America. *J. Geophys. Res.*, 86, B11, 10753-10770.
- Sajona, F. G., Maury, R. C., Bellon, H., Cotten, J., Defant, M. J., Pubellier, M. and Rangin, C., 1993. Initiation of subduction and the generation of slab melts in western and eastern Mindanao, Philippines. *Geology*, 21, 1007-1010.
- Smith D. R. and Leeman, W. P., 1987. Petrogenesis of Mount St Helens dacitic magmas. *J. Geophys. Res.*, 92, B10, 10313-10334.
- Winter, T., 1990. Mécanismes des déformations récentes dans les Andes équatoriennes. Thèse Paris XI, Orsay, France, 205p.

