

FIG.3. The model velocity field in the study area on May 28, 1985. Representative current speeds (cm/sec) are shown at a few locations.

results for the upper layer velocity field in the throughflow region on May 28, 1985. The strong southward flow in the Makassar Strait ($30\text{--}60\text{ cm}\cdot\text{s}^{-1}$), fed from the Sulawesi Sea, bifurcates at the archipelago, one branch entering directly into the Indian Ocean through the Lombok Strait and the other branch flowing east through the Flores and Banda seas. The time series of transport predicted by the model for 1985-1986 is compared to our observations in Figure 4.

The computations of observed transport through the Lombok Strait presented in (Murray and Dharma Arief, 1988) are based on monthly block averages of current meter data, which allowed for temporal variation in the vertical current profile. In order to increase temporal resolution for a comparison to the model output, the carefully determined monthly average transports were regressed against monthly average velocity components from various current meters. A best-fit relation ($R^2 = 0.79$) determined over the range of the monthly averages ($5 \leq v \leq 55\text{ cm}\cdot\text{s}^{-1}$):

$$T = 0.64 e^{.032 v}$$

where T is transport in Sv and v is the monthly average north-south component of current speed at the 35-m level (v_{35}) at the site 2 mooring, gave best results. Utilizing this relationship, the v time series from this current meter was passed through a 30-day low-pass filter (v (LP)) and used to compute transport. For the limited amount of data outside the range of the monthly averages where v (LP) $< 5\text{ cm}\cdot\text{s}^{-1}$, we let $T \rightarrow 0$ linearly with v (LP). Correspondingly, when v (LP) $> 55\text{ cm}\cdot\text{s}^{-1}$, an algorithm where $v = \text{constant} = v_{35}$ in the upper 100 m and decreases parabolically to 0 at 200 m gave good results.

Figure 4 shows a comparison between transport through the Lombok Strait calculated from the observations in this manner and from the numerical model. In all figures, eastward- and northward-directed flows are positive; westward- and southward-directed flows are negative. Owing to restrictions on the model grid size, the Lombok opening in the model for the upper active layer is about 2.2 times the actual cross section of the strait above the 200-m depth level, where the net transport is calculated. Note a transport scale adjusted for this discrepancy on the figure. We note the general agreement in phase

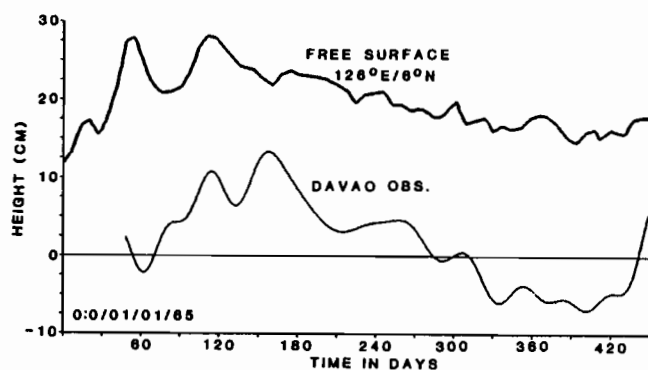


FIG.10. Observed sea level at Davao compared to sea level at the model grid point off the southern Philippines at 126°E, 6°N.

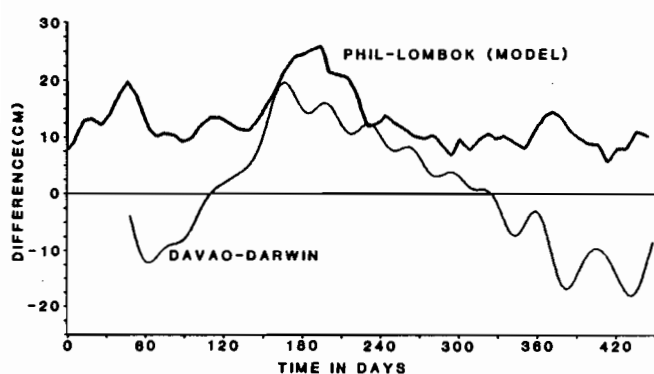


FIG.11. The Davao-Darwin sea level differences, January 1985-March 1986, compared to the sea level difference between the model grid points off the southern Philippines and the Lombok Strait.

4. Summary

Observations of currents, transport, sea level, and sea level slope in Lombok Strait and west Flores Sea in 1985 and 1986 have been compared to a simulation of the NORDA global ocean model in the Indonesian throughflow region. Despite the relatively coarse grid scale of the model compared to topographic length scales in the region, the model appears to be realistically reproducing many of the observed features. The predicted transport through the Lombok Strait, for example, agrees with detailed observations of phase and magnitude, especially when corrected for grid size limitations. The sea surface fluctuations and sea surface slopes predicted by the model agree within less than a factor of 1.5 with sea level changes and slopes observed on tide gauges. There do appear to be several cases of phase difference of several months between model and observations that require further investigation.

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**WESTERN PACIFIC INTERNATIONAL MEETING
AND WORKSHOP ON TOGA COARE**

Nouméa, New Caledonia

May 24-30, 1989

PROCEEDINGS

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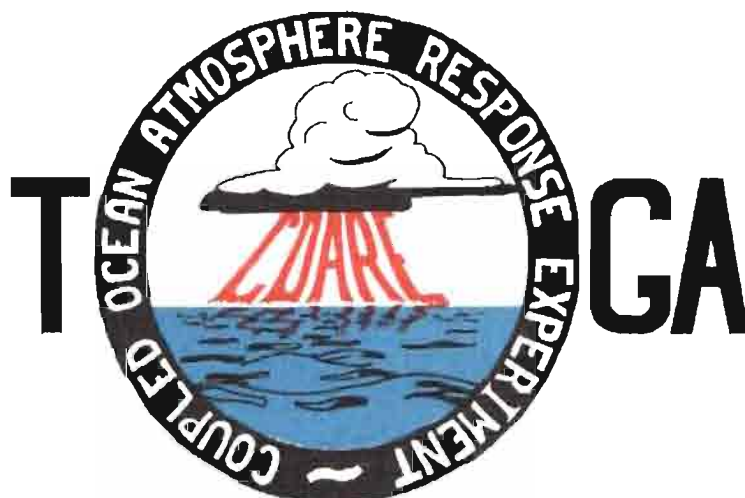


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