

## A New Ocean GCM for Tropical Ocean and ENSO Studies

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### ABSTRACT

A new reduced gravity, primitive equation model of the upper equatorial ocean has been developed by Gent and Cane (1989). It has been configured to simulate the annual cycle in the tropical Pacific Ocean. It is forced by the monthly winds from Rasmusson and Carpenter (1982) and a new heat flux formulation from Seager et al. (1988). The coefficients of vertical eddy viscosity and conductivity depend upon the Richardson number of the flow. With this mixing a deep warm pool forms in the Western Pacific, increasing the west-east gradient of the thermocline and equatorial undercurrent. The model has been diagnosed as to what factors maintain this deep warm pool.

### 1. Heat flux parameterization.

Heat Flux = Incoming Solar - Latent - (Sensible + Longwave)

\* Incoming Solar Flux =  $(1 - A) (1 - 0.62 C + 0.0019 \theta) Q_0$   
A is albedo, C is % cloud cover,  $\theta$  is solar altitude,  $Q_0$  is clear sky flux

\* Latent Heat Flux =  $\rho_a C_E L |\mathbf{u}| (1 - \delta) q_s(\text{SST})$   
 $|\mathbf{u}|$  is wind speed with minimum of  $4 \text{ m s}^{-1}$ ,  $q_s(\text{SST})$  is saturation humidity, and humidity factor  $\delta$  is related to relative humidity,  $\delta_{\text{rh}}$ ,

$$\delta q_s(\text{SST}) = \delta_{\text{rh}} q_s(T_a)$$

\* Sensible + Longwave Flux =  $\alpha (\text{SST} - T^*)$

$$[A = 0.06, \delta = 0.78, \alpha = 1.8 \text{ m s}^{-1}, T^* = 2.8^\circ\text{C}]$$

### 2. Vertical mixing coefficients.

\* Diffusivity:  $\nu = 10^3 (1 + 10R_i)^{-2} + \nu_B \quad \text{cm}^2 \text{ s}^{-1}$

\* Conductivity:  $K = 10^3 (1 + 10R_i)^{-3} + K_B \quad 0 < R_i < 3$

$\nu_B(z)$  varies between 1 and  $2 \text{ cm}^2 \text{ s}^{-1}$   
 $K_B(z)$  varies between  $1/2$  and  $1 \text{ cm}^2 \text{ s}^{-1}$

In the COARE volume,  $R_i$  is nearly always  $> 3$ , so that the vertical mixing is done solely by the background values  $\nu_B$  and  $K_B$ .

Sections of average temperature and zonal velocity in January along the equator down to 400 m and along  $160^\circ \text{ E}$  down to 500 m are shown in Figures 1-4.



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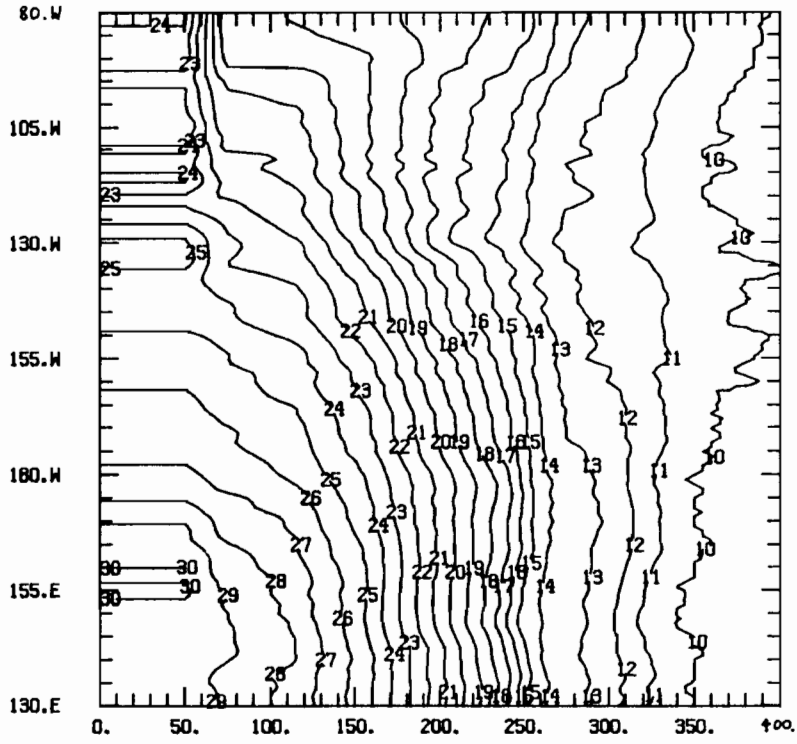


FIG.1. Average January Temperature ( $^{\circ}$ C) along Equator.

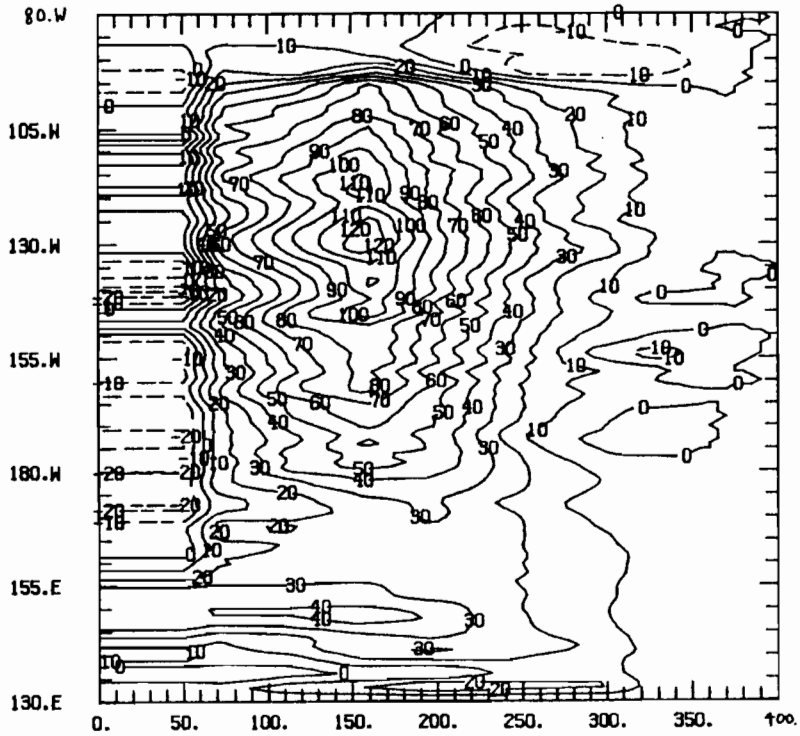


FIG.2. Average January Zonal velocity ( $\text{cm}\cdot\text{s}^{-1}$ ) along Equator.

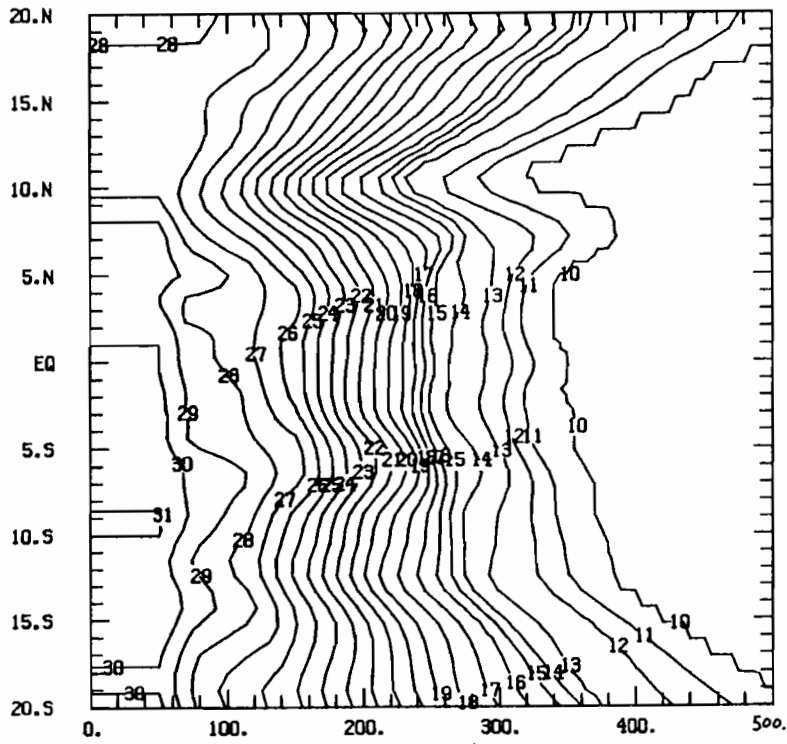


FIG.3. Average January Temperature ( $^{\circ}\text{C}$ ) along  $160^{\circ}\text{E}$ .

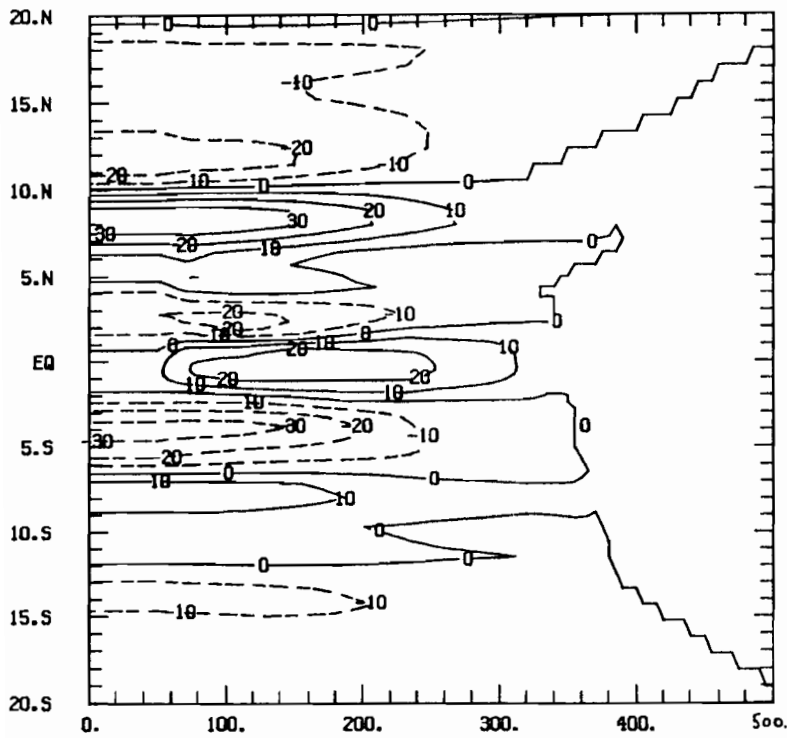


FIG.4. Average January Zonal Velocity ( $\text{cm}\cdot\text{s}^{-1}$ ) along  $160^{\circ}\text{E}$ .

### 3. Heat balance over the TOGA-COARE domain.

140°E-180°E; 10°S-10°N; 0-150 m (average)

Advection and temperature gradients in the domain are small, and the advection tends to be parallel to the isotherms, so that the 3-D heat advection in the domain is small. Thus, if the vertical heat diffusion is small, then the heat flux must also be small. This is confirmed in the model, where over the COARE domain Ri mixing and horizontal diffusion are negligible. Thus, over a year

$$\begin{array}{rclcl} \partial/\partial t(hT) = & -3\text{-D Advection} & + \text{Heat Flux} & - \text{Vertical diffusion} & \\ 0 \approx & -3.9 & + 13.5 & - 9.4 \text{ W m}^{-2} & \end{array}$$

The model SST for January and the SST such that  $Q \approx 0$  (Figures 5 and 6) show that, over the COARE area, SST is set by  $Q(\text{SST}) \approx 0$ .

$$\begin{array}{rclcl} \text{Heat Flux} = & \text{Incoming Solar} & - \text{Latent} & - (\text{Sensible} + \text{Longwave}) & \\ 13,5 = & 189,8 & - 117,6 & - 58,7 \text{ W m}^{-2} & \end{array}$$

$$\begin{array}{rclcl} \nabla \cdot (huT) & = & \partial/\partial x(huT) & + \partial/\partial y(hvT) & + \partial/\partial s(wT) \\ 3,9 & = & -246,3 & + 202,1 & + 48,1 \text{ W m}^{-2} \end{array}$$

Thus the 3-D heat advection is a small residual of large component terms and, therefore, will be very difficult to measure. Alternatively, a small error in the very small angle between velocity and isotherms will cause a large error in the 3-D advective heat flux estimate. Quantitatively

$$|\nabla \cdot (huT)| / \{ \sum [\partial/\partial x_i(hu_iT)]^2 \}^{1/2} = 0.012$$

### 4. What sets the SST and depth of the warm pool?

The hypothesis that  $0 < Q(\text{SST}) < 20 \text{ W m}^{-2}$  was confirmed by setting the minimum value of  $|u|$  to  $3 \text{ m s}^{-1}$  (not 4) in the latent flux. This increases the temperature, such that  $Q = 0$ , from about  $29.5 \rightarrow 30^\circ\text{C}$  (min = 4) to about  $33 \rightarrow 33.5^\circ\text{C}$  (min = 3) on the equator between  $130^\circ\text{E}$  and  $150^\circ\text{E}$ . The actual model SST in this region increases from about  $29.5^\circ\text{C}$  (min = 4) to about  $32^\circ\text{C}$  (min = 3). When the minimum is set to  $3 \text{ m.s}^{-1}$ , the outgoing latent flux is reduced by 1/4 in the region, or about  $30 \text{ W m}^{-2}$ . Thus, the sensitivity of SST in this region to errors in the heat flux is

$$\text{Sensitivity} \equiv dT/dQ = 2.5/30 = 1/12 \text{ }^\circ\text{C W}^{-1} \text{ m}^{-2}$$

Thus, an error of  $16 \text{ W m}^{-2}$  in  $Q$  will lead to an error of  $1^\circ\text{C}$  in SST in this region of the model. With  $|u| \text{ min} = 3 \text{ m s}^{-1}$ , the heat balance over a year is

$$0 \approx -5.5 (3 - D A) + 17.7 (HF) - 10.5 (VD) \text{ W m}^{-2}$$

This similar to the min =  $4 \text{ m s}^{-1}$  case, but is not so close to equilibrium.

The depth of the warm pool in this model is set by  $K_B$ . In a run where  $K_B$  and  $v_B$  were increased by factors of 2 and 10, with no  $R_i$  mixing, the depth of the warm pool (water  $> 27^\circ\text{C}$ ) was reduced by about 50m. Thus the run with smaller  $K_B$ , which can sustain a sharper thermocline, has a deeper warm pool. With a larger  $K_B$ , there is more vertical diffusion of heat, so that the COARE region SST is reduced by close to  $1^\circ\text{C}$  to produce a larger heat flux. With increased  $v_B$  and  $K_B$ , the heat balance over a year is

$$0 \approx 3.2 (3 - D A) + 25.0 (HF) - 21.3 (VD) \text{ W m}^{-2}$$

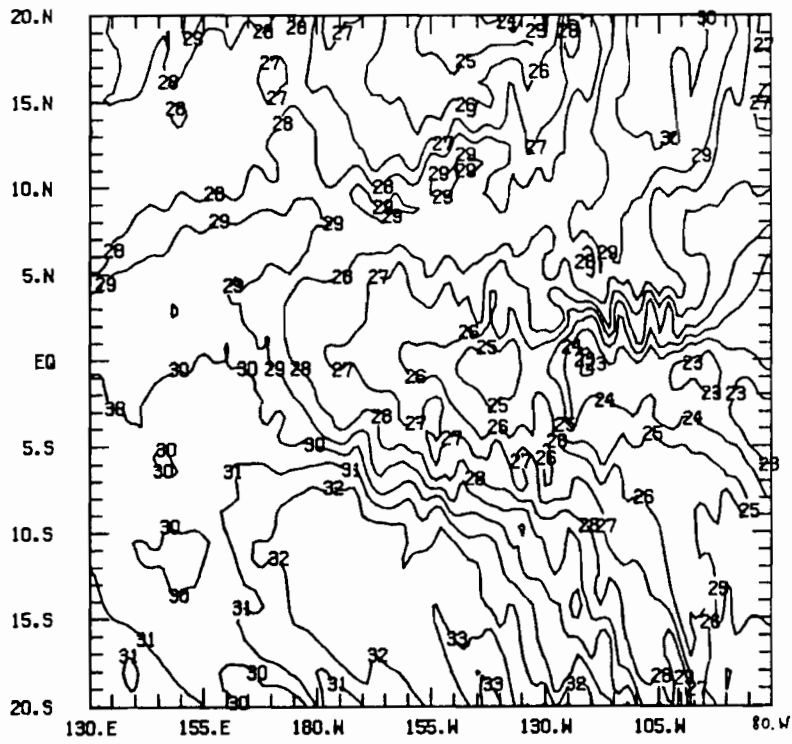


FIG.5. Average January Sea Surface Temperature (°C).

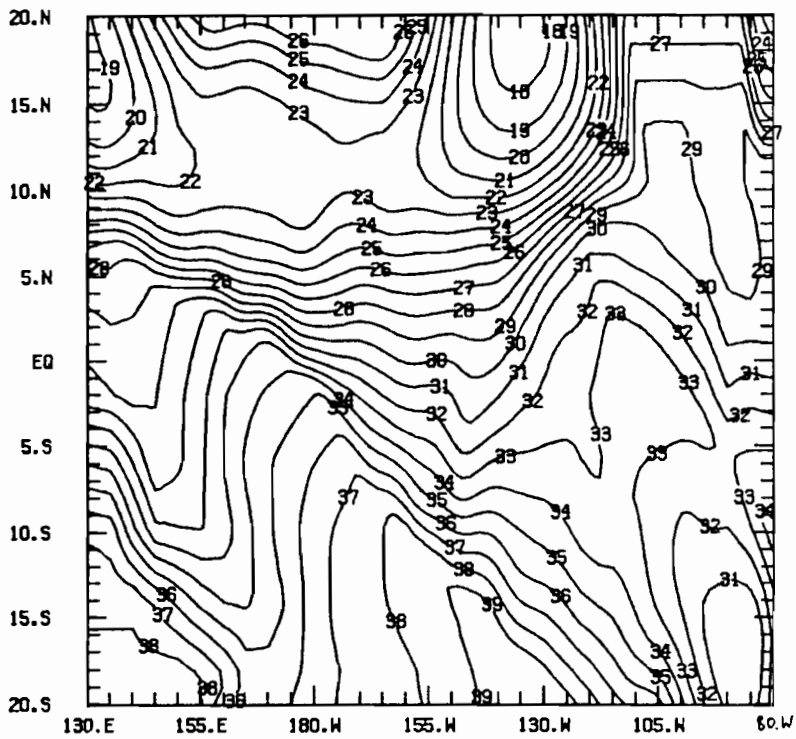


FIG.6. Average January Temperature (°C) such that  $Q = 0$ .

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**WESTERN PACIFIC INTERNATIONAL MEETING  
AND WORKSHOP ON TOGA COARE**

**Nouméa, New Caledonia**

**May 24-30, 1989**

**PROCEEDINGS**

*edited by*

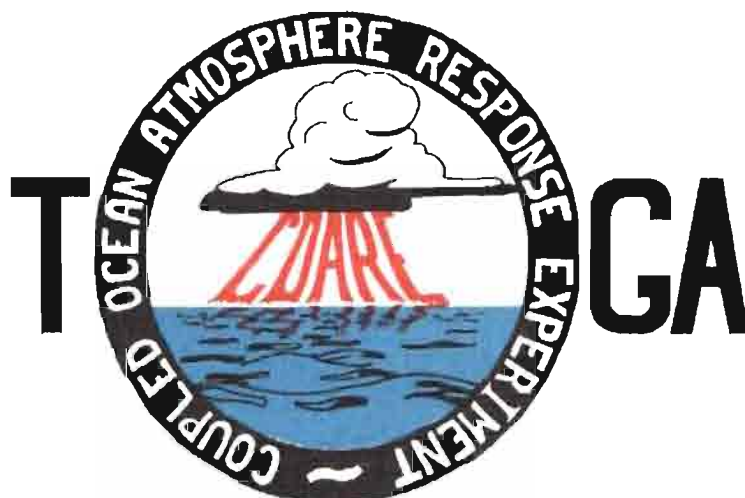
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