

OVERGRAZING EFFECTS ON SOIL PROPERTIES IN THE PÁRAMO OF LLANGAHUA AND ESPERANZA (TUNGURAHUA, ECUADOR)

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High Andean grassland, or "Páramos", is a high-altitude ecosystem. It is usually located between the upper part of the mountain forest at 3500m above sea level and the lower part of the permanent snow at 4500 m a.s.l. The Páramo occurs irregularly in Venezuela, Colombia and Ecuador. The continuous grass layer forms the dominant ground cover. Soils developed on recent (Holocene) volcanic ashes in the entire northern half of Ecuador. The high water-retention properties of these soils associated with the specific vegetation confer to the Páramo an important role in water availability which is paramount for the majority of the Ecuadorian population (Luteyn, 1992). The study area is located in the province of Tungurahua on the occidental chain, 20 km north-western of the basis of the Chimborazo volcano with an elevation ranging from 3800 to 4200 m (Fig. 1). Due to a rapid increase of the rural population, especially after the agrarian reform of 1970, these areas were often degraded due to overgrazing. This study intended to show the effects of overgrazing by sheep on soil properties and vegetation in a transition zone from a sub-humid to a dry Páramo, managed by an indigenous community.

Soil properties. Due to the homogeneity of the recent deposits, most probably originated from the same source (Cotopaxi and/or Tungurahua; Winckell et al., 1991), the pedological differentiation results only from different climatic conditions (Colmet-Daage et al, 1967.). In the humid zone (north of the studied area), the soils have an udic regime and generally andic properties with a low bulk density (< 0.9), high accumulations of organic matter ($> 6\%$), and contain amorphous constituents (Quantin, 1995; Shoji et al, 1996). In the dry zone soils have an ustic regime with vitric properties, exposing a high amount ($> 60\%$) of non-weathered primary minerals. Both types of soils have a bimodal granulometric composition (peaks at $10 \mu\text{m}$ and $100 \mu\text{m}$) typical of volcanic deposits while in the southern part of the area soils are highly aeolian dependent with a monomodal granulometric peak at $100 \mu\text{m}$ which leads to a rapid transition to a xeric regime.

Overgrazing effects: Two main overgrazing effects impact on vegetation. Firstly, a strong decrease of the biodiversity (Jørgensen and Ulloa, 1994) and the accompanying biomass within the same area. This was similar to the effect of climatic change as evidenced in a more extended zone (Hofstede et al, 1995 ; Pels and Verweij, 1992). Secondly, the increase of the bare surface due to trampling which in turn increases the risk of erosion by favouring preferential paths for water flows and consecutive of steep slopes.

Destruction of the soil-cover results in the direct exposure of the topsoil at the surface and eventually to irreversible desiccation. This degradation appears as a retrogradation of evolved vitrisols-andisols to weakly evolved vitrisol by an increasing content of sand and slightly weathered minerals in the surface horizons, and, also to a decrease of the CEC. The $\text{Al}+1/2\text{Fe}$ oxalate extract decreases from about 1.2% to less than 0.3%. The bimodal granulometric composition is less apparent in the upper horizon with a decrease of the $10 \mu\text{m}$ particles content. A possible segregation of fine particles is possible by lack of grass cover. This retrogradation is accompanied with a decrease of the Carbon content, most likely by mechanical erosion, from about 10g/kg to 4 g/kg. Moreover, the degradation of the soil cover induces a serious decrease of the soil-water content. In a humid area the soil moisture decreases from 70% to 15 % and in a dry area from 30% to less than 10%. However, it shows an important decrease of the water content at pF 2.5 after a complete saturation during 36 hours, at least. This means that the soil developed a strong water repellency which is more significant below the tussocks than on the bared surface.

Degradation of the Páramos by overgrazing does not affect solely the biodiversity, but also the soil properties. It results in a severe decrease of the Carbon content and an irreversible decrease of the

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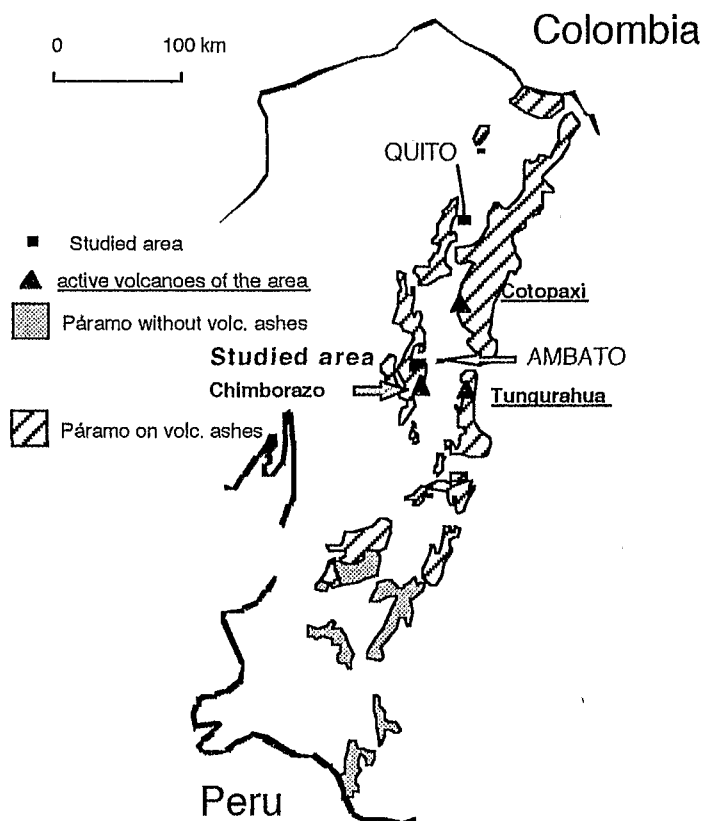
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soil-water content. The main recourse of sheep breeders is the selling of organic manure to the market gardeners of the Ambato valley so that the organic matter deficit is incremented. The organic layer of degraded soils exhibits also a strong water repellence which increases the erosion sensibility and limits water infiltration. In a long term perspective, the water reserves of the Páramo are expected to decline..

Fig. 1 : Location of the studied area



References :

- Colmet-Daage, F., Cucalon, F., Delaune, M., Gautheyrou, J., Gautheyrou, M., et Moreau, B., 1967. Caractéristiques de quelques sols d'Equateur dérivés de cendres volcaniques. 2ème partie : Conditions de formation et d'évolution. Cah. ORSTOM, sér. pédol., Vol. V, n°4: 353-392.
- Hofstede, R.G.M., Mondragon, M.X. and Rocha, C.M., 1995. Biomass of grazed, burned and undisturbed páramo grasslands, Colombia.I. Above ground vegetation. Arctic and alpine Research 27 (1): 1-12.
- Jørgensen, P.M. and Ulloa Ulloa, C., 1994. Seed plants of the high Andes of Ecuador. A checklist. Department of Systematic Botany, University of Aarhus. AAU reports 34. Aarhus, Denmark, 445 p.
- Luteyn J. L., 1992 - Paramos : why study them ? *In* Paramo : An andean ecosystem under human influence. H. Baslev and J. L. Luteyn, Academic Press, London, 1-14.
- Pels, B. and Verweij, P.A., 1992. Burning and grazing in a bunchgrass páramo ecosystem: Vegetation dynamics described by a transition model. Burning and grazing gradients in páramo vegetation: Initial ordination analyses. In : Balslev, H. and Luteyn, J.L. eds. Páramo. An andean ecosystem under human influence. Academic Press, 177-195.
- Quantin, P., 1995. Andosols et vitrosols. Référentiel pédologique français. INRA - AFES, 85-93.
- Shoji, S, Nanzyo, M., Dahlgren, R.A. and Quantin, P., 1996. Evaluation and proposal revisions of criteria for andosols in the world reference base for soil resource. Soil Science, vol.161, n°9 : 604-615.
- Winckell A., Zebrowski C. et Delaune M., 1991. Evolution du modèle quaternaire et des formations

Table 1 : some chemical properties of the upper horizons of characteristic profiles of Llangahua and La Esperanza

profiles	horizons	pH H ₂ O	pH KCl	C g/100g	N g/100g	C/N	Exchangeable complex							
							Ca ⁺⁺	Mg ⁺⁺	K ⁺	Na ⁺	Al ³⁺	S	CEC (T)	S/T %
HUMID PARAMO														
not degraded	LLA 20.1.	6.9	6.5	11.55	0.951	12.1	37.90	5.27	1.95	0.37	0.06	45.49	45.59	100
low flat area	LLA 20.2.	7.0	6.6	8.01	0.532	15.1	29.80	9.72	1.29	0.32	0.05	41.13	41.20	100
hill crest	ESP 41.1.	4.8	4.8	10.11	0.624	16.2	18.10	3.30	0.49	0.38	0.16	22.27	22.54	99
	ESP 41.2.	4.7	4.7	6.82	0.478	14.3	11.30	1.98	0.20	0.19	0.35	13.67	14.08	97
recolonization	LLA 21.1.	5.4	5.0	9.94	0.690	14.4	23.20	5.82	0.47	0.34	0.10	29.83	29.98	100
	LLA 21.2.	5.4	4.8	8.76	0.606	14.5	21.00	5.12	0.37	0.42	0.16	26.91	27.14	99
degraded weakly	ESP 61.1.	5.2	5.1	7.94	0.495	16.0	22.90	6.41	0.41	0.34	0.10	30.06	30.18	100
	ESP 61.2.	5.0	4.7	6.47	0.408	15.9	17.10	5.62	0.26	0.41	0.15	23.39	23.55	99
	LLA 22.1.	5.9	5.2	8.54	0.577	14.8	24.10	5.99	0.48	0.31	0.10	30.83	30.95	100
	LLA 22.2.	5.9	5.1	7.48	0.502	14.9	21.60	5.62	0.39	0.50	0.12	28.11	28.24	100
degraded strongly	LLA 23.1.	5.4	4.9	4.74	0.322	14.7	18.30	5.98	0.37	0.61	0.13	25.26	25.42	99
	LLA 23.2.	5.4	5.4	3.53	0.239	14.8	16.80	5.53	0.32	0.65	0.03	23.30	23.33	100
below tussock	LLA 24.1.			7.65	0.547	14.0	19.30	3.46	0.44	0.29	0.22	23.49	23.78	99
	LLA 24.2.			6.10	0.516	11.8	21.90	6.69	0.26	0.41	0.14	29.26	29.42	99
beside tussock	LLA 25.1.			5.56	0.444	12.5	20.00	5.42	0.39	0.31	0.10	26.06	26.18	100
	LLA 25.2.			4.66	0.347	13.4	18.80	5.25	0.36	0.40	0.10	24.81	24.92	100
DRY PARAMO														
not degraded	ESP 31.1.	5.3	5.3	6.95	0.425	16.4	20.50	4.96	0.61	0.21	0.08	26.28	26.37	100
	ESP 31.2.	5.0	5.0	5.44	0.380	14.3	16.10	4.39	0.27	0.38	0.11	21.14	21.27	99
degraded flat	ESP 21.1.	5.4	5.8	4.22	0.340	12.4	15.30	2.00	0.27	0.23	0.04	17.80	17.84	100
	ESP 21.2.	5.6	6.0	1.74	0.161	10.8	12.40	1.63	0.17	0.21	0.01	14.41	14.42	100
degraded slope below tussock	ESP 11.1.	4.8	5.1	4.02	0.280	14.3	11.90	2.27	0.33	0.16	0.06	14.66	14.73	100
	ESP 11.2.	5.7	6.0	1.70	0.133	12.8	13.50	3.17	0.52	0.23	0.02	17.42	17.44	100
EL ARENAL														
below tussock	ESP 1.2.	5.3	5.3	0.70	0.063	11.1	3.93	0.95	0.11	0.10	0.08	5.09	5.17	98
	ESP 1.3.	5.6	5.8	0.79	0.064	12.4	10.40	1.91	0.40	0.25	0.04	12.96	13.00	100
PANTANO														
organic matter	ESP 51.1	5.7	5.2											
	ESP 51.2.	5.2	5.1	12.42	0.966	12.9	23.40	8.36	0.57	0.52	0.08	32.85	33.00	100
	ESP 51.3.	5.3	5.1	7.52	0.559	13.5	20.10	6.42	0.42	0.57	0.07	27.51	27.62	100

Table 2. Some physical properties of the upper horizons of some characteristic profiles of Llangahua and La Esperanza

profiles	horizons	text ure						moisture g.100g ⁻¹	pF 300 KPa g.100g ⁻¹	B.D.
		CSa	Fsa	CSi	FSi	Cl	H ₂ O			
HUMID PARAMO										
not degraded	LLA 20.1.	3.1	15.0	2.1	34.0	35.5	5.9	53.08	68.37	
low flat area	LLA 20.2.	4.7	23.7	2.4	8.2	52.6	4.5	40.81	51.69	
hill crest	ESP 41.1.	3.7	22.3	10.5	33.9	6.9	4.3	89.80	77.61	0.68
	ESP 41.2.	5.7	26.4	7.6	30.4	12.7	4.6		70.48	
recolonization	LLA 21.1.	3.7	21.6	5.8	18.7	43.3	4.6	48.76	54.02	0.78
	LLA 21.2.	4.1	23.2	6.0	29.9	28.9	4.3	49.85	52.53	
degraded weakly	ESP 61.1.	3.4	16.5	9.8	30.6	3.3	5.1	31.30	42.59	0.80
	ESP 61.2.	4.5	24.6	9.1	32.5	10.4	4.0	24.60	44.48	0.88
	LLA 22.1.	4.2	23.4	2.2	16.7	43.5	4.1	25.77	42.86	0.78
	LLA 22.2.	5.1	24.8	4.7	14.3	42.7	4.1	25.38	42.20	
degraded strongly	LLA 23.1.	6.7	29.4	3.5	11.4	41.2	3.0	10.89	19.80	0.76
	LLA 23.2.	8.1	30.1	9.0	27.7	17.9	2.7	15.47	25.26	
below tussock	LLA 24.1.	5.6	29.8	8.3	27.1	10.5	3.3	9.09	17.41	
	LLA 24.2.	5.8	27.8	6.7	18.5	24.7	3.2	11.19	25.54	
beside tussock	LLA 25.1.	6.7	29.4	8.8	27.8	10.9	3.1	11.31	28.88	
	LLA 25.2.	7.8	30.4	6.8	28.5	11.8	2.9	12.45	30.55	
DRY PARAMO										
not degraded	ESP 31.1.	4.3	26.0	14.0	28.8	8.1	3.8	31.31	31.54	0.74
	ESP 31.2.	6.1	28.4	3.6	29.4	19.7	3.5	30.89	32.93	0.93
degraded flat	ESP 21.1.	8.0	36.2	13.9	26.2	6.3	1.8	10.43	15.54	0.99
	ESP 21.2.	10.9	37.7	9.2	31.2	5.5	1.3		14.04	
degraded slope below tussock	ESP 11.1.	5.0	46.7	14.3	20.6	3.2	1.7	8.10	9.45	0.83
	ESP 11.2.	10.9	36.1	7.5	33.7	8.7	1.7	5.90	8.75	1.12
beside tussock	ESP 12.1							11.22		1.11
EL ARENAL										
below tussock	ESP 1.2.	18.5	59.8	3.8	10.4	4.7	0.8	7.99	16.29	1.21
	ESP 1.3.	15.0	49.0	6.8	19.1	6.5	1.9	18.96	25.72	1.30
beside Tussock	ESP 2.1							18.40		1.28
PANTANO										
organic matter	ESP 51.1.							472.71	399.26	0.11
	ESP 51.2.	5.2	20.5	10.7	30.8	8.2	4.8	206.25	150.4	0.36
	ESP 51.3.	5.3	27.6	7.1	31.6	13.6	4.4		65.07	

Csa : Coarse sands ; Fsa : fine sands ; Csi : coarse silts ; Fsi : fine silts ; Cl : clay ; B.D. : bulk density

Table 3. Extraction of amorphous components of the upper horizons of some characteristic profiles of Liangahua and La Esperanza

profiles	horizons	oxalate				pyrophosphate			
		Si	Fe	Al	Al+1/2 Fe	Si	Fe	Al	Alp/Alo
		g.kg ⁻¹				g.kg ⁻¹			
HUMID PARAMO									
not degraded	ESP 41.2.	2.30	7.40	9.10	1.28	0.87	4.51	5.30	0.58
degraded	ESP 61.2.	1.39	5.34	5.80	0.85	0.70	3.38	3.61	0.62
DRY PARAMO									
not degraded	ESP 31.2.	0.93	3.94	4.00	0.60	0.57	2.36	2.57	0.64
degraded flat	ESP 21.2.	0.72	1.63	1.94	0.28	0.23	0.27	0.60	0.31
degraded slope	ESP 11.2.	0.54	1.87	1.62	0.26	0.31	0.35	0.40	0.25
EL ARENAL									
below tussock	ESP 1.2.	1.11	1.83	2.36	0.33	0.21	0.28	0.56	0.24