

SEASONAL ESTIMATES OF RIPARIAN EVAPOTRANSPIRATION  
(CONSUMPTIVE WATER USE) USING REMOTE AND *IN-SITU* MEASUREMENTS

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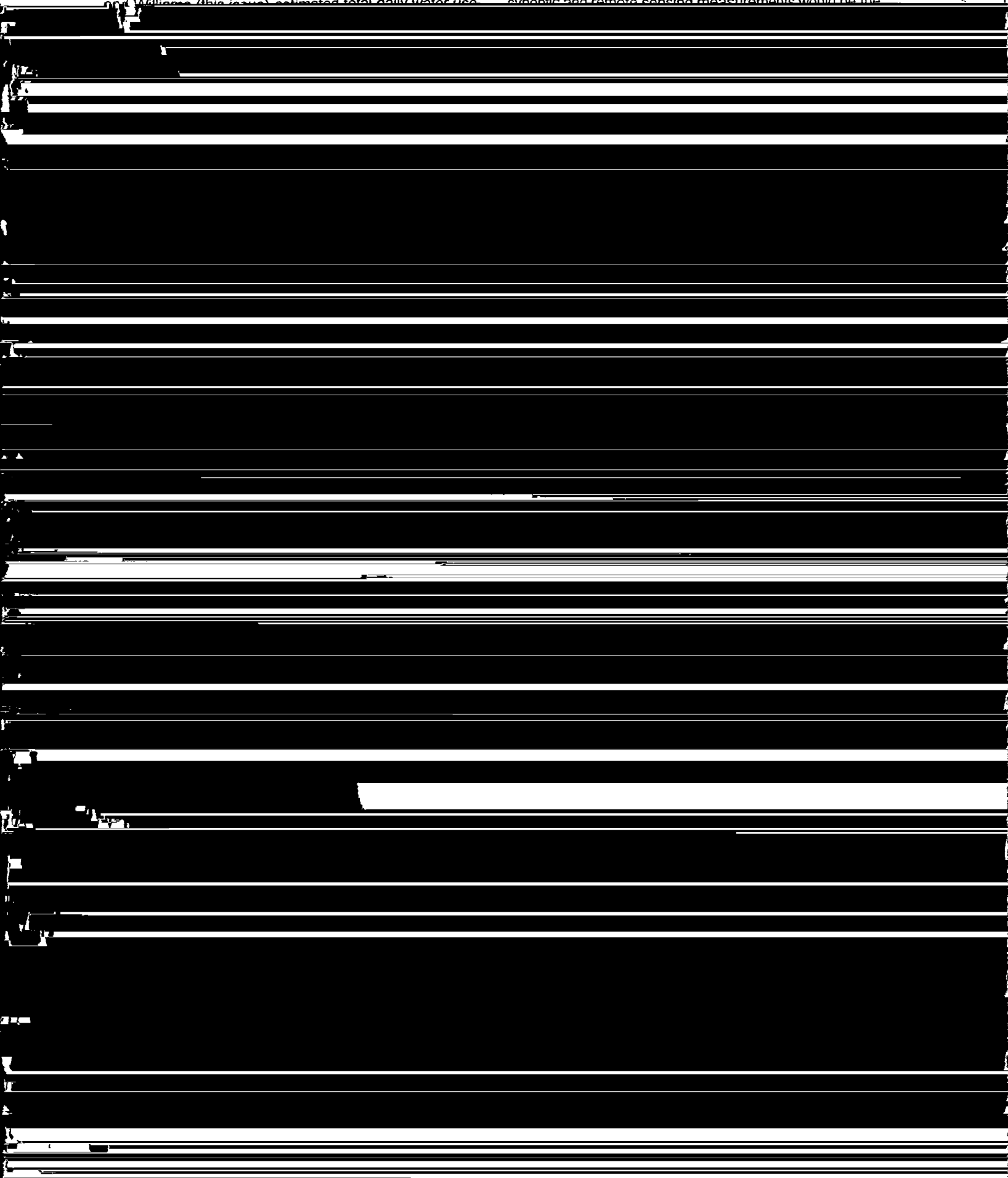
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not Williams (this issue) estimated total daily water use synoptic and remote sensing measurements would be the



Charleston Gage for the 1996 water year. The shaded portion of the figure is an estimate of the riparian ET that abstracts water from all baseflow stream inputs. For this example it was assumed that steady-state baseflow unaffected by riparian ET was equal to be the average winter discharge. During monsoon storm flows the baseflow was linearly interpolated between June 28 and Oct. 1. Using these assumption the riparian ET is 9,928,800 m<sup>3</sup> (3,185 ac-ft). Although this method is straightforward it invokes several strong assumptions. First that groundwater inflow is steady throughout the river reach of interest. In the San Pedro this assumption is tenuous at best given the significant pumping taking place in the basin. Secondly it assumes baseflows can be easily separated from storm runoff. As noted by Maddock et al. (this issue), this is not a simple task. Finally, this method is only applicable in perennial stream reaches. If a reach becomes intermittent, as occurred on the San Pedro in 1997 downstream of the Charleston gage, this type of analysis cannot be carried out. This also points to another problem in this analysis. A related problem is that there is uncertainty in the length of river reach over which riparian ET losses are reflected in the downstream discharge measurements.

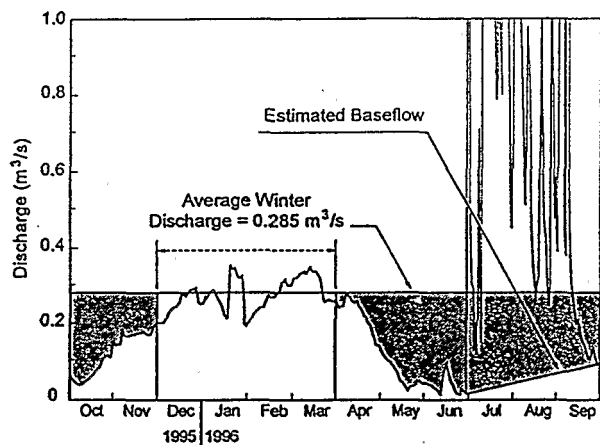


Figure 2. San Pedro River discharge at Charleston, AZ for the 1996 water year.

Other atmospheric controls are also important for estimation of riparian ET. Air temperature, as noted above, is highly correlated to radiation or the available energy for driving ET. However, temperature-based ET-estimation methods are not recommended unless this is the only available data source (Shuttleworth, 1992). Shuttleworth (1992) also notes that preferred methods for estimating ET require a value of the difference between the saturated vapor pressure ( $e_s$ ) and the ambient vapor pressure ( $e$ ), or vapor pressure deficit ( $VPD = e_s - e$ ). VPD and whole-tree transpiration (tree #203) are plotted for several days of the August campaign in Figure 3. A reasonable correlation exists between these quantities indicating a simple model that incorporates VPD could be utilized to estimate the temporal variations in ET between intensive measurement campaigns where remote sensing data was acquired. Variable cloudiness can complicate

the sap flow versus VPD relationship. It can be improved if it is constrained to only well lit (high radiation:  $> 75 \text{ W/m}^2$ ) conditions.

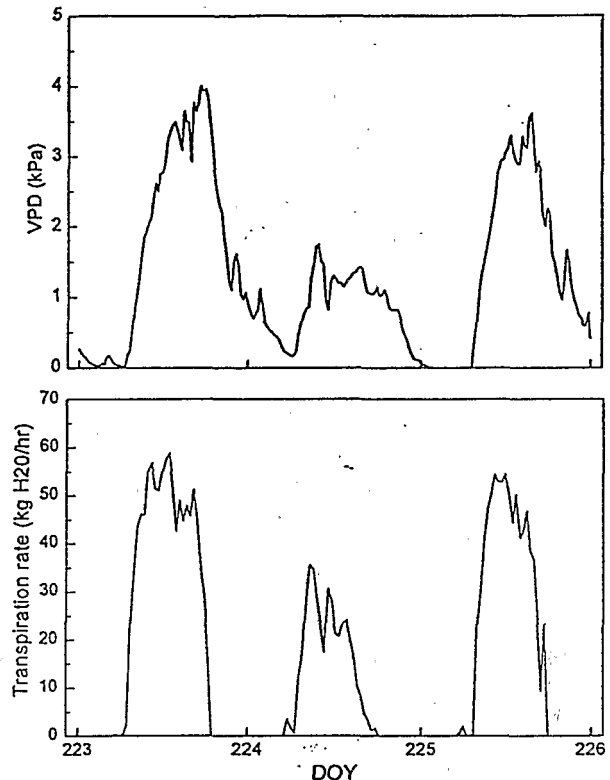


Figure 3. Vapor Pressure Deficit (VPD) and transpiration from a large cottonwood for Julian days 223-225 (August 11-13).

Williams et al. (this issue) and Schaeffer and Williams (this issue) noted the strong influence of photosynthetically active radiation (PAR) on the diurnal variations in whole tree transpiration. Their plots of PAR versus whole tree transpiration for a portion of the same time shown in Figure 3 are even more clearly correlated than the VPD plot. Note that air temperature and VPD form the basis for the calculation of potential evapotranspiration which will also be considered in future analysis.

### 3. CONCLUSIONS - FUTURE DIRECTIONS

Additional variables will be examined to better estimate the temporal variation in riparian ET between intensive measurement periods as well as future work with models. Various space-time interpolation treatments for seasonal changes of the inferred differential riparian stress in space obtained from the multi-temporal remotely sensed data will also be examined. With this information it is expected that reasonable estimates of the seasonal variation in riparian ET as well as the total consumptive water use can be obtained in the near future.

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