

Eburnean deformation pattern of Burkina Faso and the tectonic significance of shear zones in the West African craton

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Abstract – Shear zones of the Paleoproterozoic Eburnean accretionary Orogen (West African craton) are investigated by means of large-scale structural mapping. Regional scale (10-100 km) mapping was based on the aeromagnetic survey of Burkina Faso and craton-scale (1000 km) mapping on a compilation of fabric data. At both scales, shear zones are arranged as an anastomosed transpressional network that accommodated distributed shortening and lateral flow of the orogenic lithosphere between the converging Kénéma-Man and Congo Archean provinces. Structural interference patterns at both scales were due to three-dimensional partitioning of progressive transpressional deformation and interactions among shear zones that absorbed heterogeneities in the regional flow patterns while maintaining the connectivity of the shear zone network. Such orogen-scale kinematic patterns call for caution in using the deformation phase approach without considering the “bigger structural picture” and interpreting displacement history of individual shear zones in terms of plate kinematics. The West African shear zone pattern is linked to that of the Guiana shield through a new transatlantic correlation to produce an integrated kinematic model of the Eburnean-Transamazonian orogen.

Keywords: hot orogen / shear zone pattern / Eburnean / Transamazonian / aeromagnetic data

Résumé – **Champ de déformation éburnéen du Burkina Faso et signification tectonique des zones de cisaillement du craton ouest africain.** Une cartographie structurale à grande échelle a été entreprise dans le but d'étudier les zones de cisaillement de l'orogène éburnéen (paléoprotérozoïque) du craton ouest africain. La cartographie d'échelle régionale (10–100 km) a été réalisée à l'échelle du Burkina Faso grâce à l'interprétation d'un relevé aéromagnétique et la cartographie d'échelle cratonique (1000 km) à partir d'une compilation de données structurales. Aux deux échelles, les zones de cisaillement définissent un réseau anastomosé transpressif ayant accommodé le raccourcissement distribué et le fluage latéral de la lithosphère orogénique entre deux provinces archéennes convergentes (Kénéma-Man et Congo). Aux deux échelles, les structures d'interférence résultent d'un partitionnement tridimensionnel progressif de la déformation transpressive et des interactions entre zones de cisaillement qui ont absorbé les hétérogénéités générées dans le fluage régional tout en maintenant la connectivité du réseau de cisaillements. Un tel schéma cinématique à l'échelle de l'orogène appelle à la prudence dans l'usage des phases de déformation sans considérer la déformation à très grande échelle et l'interprétation des zones de cisaillement en termes de cinématique de plaques. Le réseau de zones de cisaillement ouest africain est repositionné vis-à-vis de celui du craton guyanais via une nouvelle reconstruction transatlantique pour produire un modèle cinématique intégré de l'orogène éburnéen-transamazonien.

Mots clés : orogène chaud / réseau de zones de cisaillement / éburnéen / transamazonien / données aéromagnétiques

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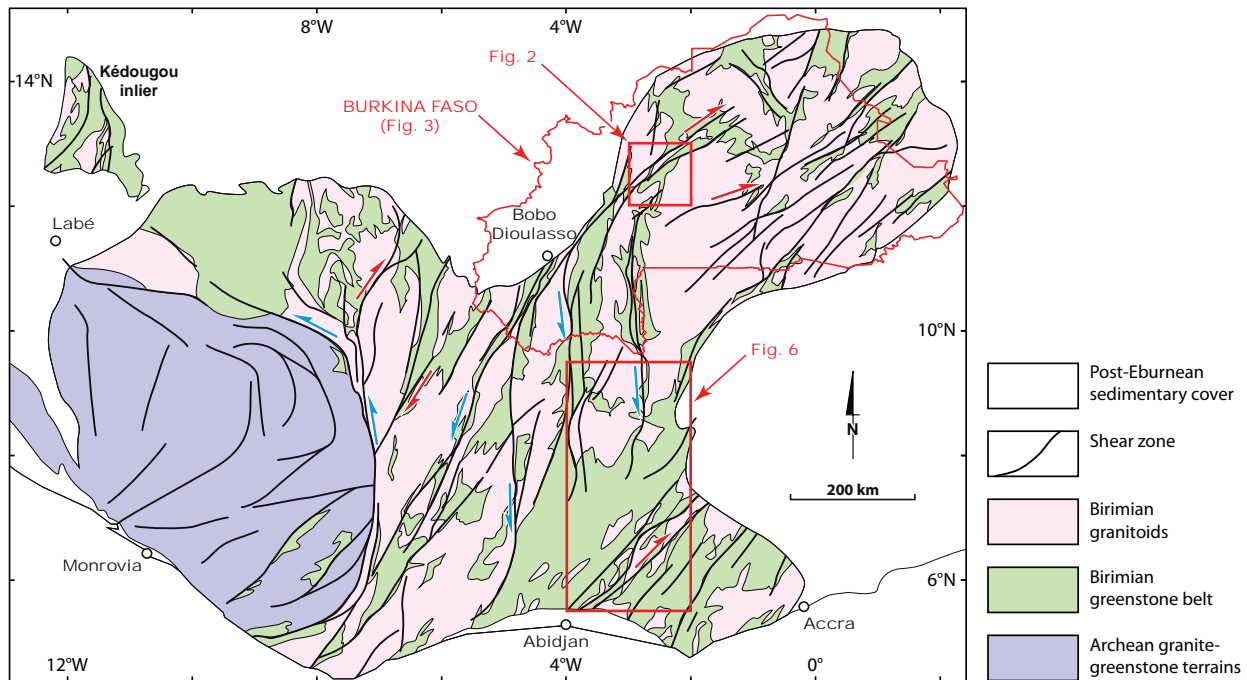


Fig. 1. Sketch structural map of the southern West African craton (after Milési *et al.*, 1989 and Thiéblemont, 2016). In the original model of Milési *et al.* (1989), the dextral shear zones (red kinematic arrows) are the latest and related to a regional D3 strike-slip deformation event, whereas sinistral ones (blue kinematic arrows) would be related to a D2 strike-slip regional deformation event.

1 Introduction

The record and geodynamic meaning of tectonic strain in Precambrian accretionary orogens are matters of renewed interest and debate. Main issues are the nature of the mountain building processes and whether all or part of those processes are the signature of modern-like plate tectonics. In absence of crustal-scale thrusts such as those documented in modern collision belts, the ubiquitous transpressive shear zones of Precambrian accretionary orogens (Cagnard *et al.*, 2006a; Chardon *et al.*, 2009; Duguet *et al.*, 2009) are central to the debate. In particular, whether they record large displacement or could be tight to a far-field kinematic framework or to a kind of plate kinematics is not straightforward (Treloar and Blenkinsop, 1995; Hammer *et al.*, 1997; Gapais *et al.*, 2005; Chardon *et al.*, 2008; Harris and Bédard, 2014). Those shear zones also tend to control the location of endogenous mineral deposits as typified by orogenic gold (e.g., Bedeaux *et al.*, 2017; Gapais, 2018). Characterization of shear zones geometries and kinematics at various scales should therefore be useful to link local deformation histories with orogen-scale processes and mineral systems.

In this contribution, we address the relationships among deformation, shear zone patterns and kinematic boundary conditions of Precambrian accretionary orogens from the regional (100 km) to the orogen (1000 km) scale. Our study is based on the example of the Paleoproterozoic Eburnean orogen, one of the largest accretionary orogens preserved on Earth. The study focuses on the largest outcrop of the orogen found in the southern West African craton (sWAC; Fig. 1). We produced a structural map at the scale of Burkina Faso

(ca. 500 × 400 km) by interpreting aeromagnetic data (Betts *et al.*, 2007), and a craton-scale map (ca. 1200 × 800 km) obtained by compiling fabric data from geological maps. Finally, a new transatlantic correlation is proposed between the sWAC and the Guiana shield to produce a structural map on a mega scale (ca. 2000 × 6000 km) and evaluate the first-order structures and kinematics of the Eburnean-Transamazonian orogen. Our results show how regional shear zone patterns can be used as a gauge of orogen-scale deformation.

2 Outline of the geological-structural context

The Eburnean orogen of the sWAC affects the so-called Birimian crust, which consists of tholeiitic to calc-alkaline mafic volcanics (preserved in greenstone belts that also comprise clastic sediments) and broadly coeval and younger TTG-like complexes and granitoid plutons (Milési *et al.*, 1989, 1992; Block *et al.*, 2016a; Parra-Avila *et al.*, 2017, 2018; Wane *et al.*, 2018; Grenholm *et al.*, 2019). Birimian mafic and felsic magmatism and Eburnean deformation and metamorphism took place essentially between 2.27 and 1.96 Ga (Grenholm *et al.*, 2019). The Birimian crust is dominantly juvenile (Abouchami *et al.*, 1990; Boher *et al.*, 1992) with an increasingly documented Pre-Birimian Paleoproterozoic and Archean components of crustal inheritance, especially approaching its bounding Archean terrains (Eglinger *et al.*, 2017; Parra-Avila *et al.*, 2018; Petersson *et al.*, 2018). The sWAC also exposes Archean terrains (in the Kénéma-Man domain), which have been involved in the Eburnean orogeny (Feybesse and Milesi, 1994; Fig. 1).

The original regional-scale kinematic model of the Eburnean orogen favors late, two-phase strike-slip-dominated deformation through the shear zone network shown on Figure 1 (Milési *et al.*, 1989; Feybesse and Milesi, 1994; Ledru *et al.*, 1994; Lompo, 2010). Pre-strike-slip deformation inferred by most authors is a matter of debate, as usually in granite-greenstone terrains. Steep fabrics in greenstone belts were interpreted as:

- originally shallow fold and thrust fabrics that were steepened and/or transposed by strike-slip deformation (e.g., Feybesse and Milesi, 1994);
- resulting from vertical tectonism due to diapirism (Vidal *et al.*, 2009; Lompo 2010).

Specific greenstone lithologies have been shown to preserve relics of HP-LT metamorphism, whereas regional HT-LP metamorphism associated with partial melting and felsic plutonism took place mostly during strike-slip dominated deformation (Ganne *et al.*, 2014). Over the last decade, the “phase” or “event” structural approach was adopted by most authors using fabric, fold or shear orientations and their relative chronology to infer successive horizontal shortening (or stretching) directions (review in Grenholm *et al.*, 2019). The tectonic significance of such directions is not straightforward, especially while attempting regional correlations amongst “events”, which are implicitly taken as indicating changes in far-field convergence (or divergence) direction at the boundaries of the orogen. The large-scale mapping approach proposed here will provide an opportunity to replace such studies in a broader structural perspective.

3 Deformation pattern at the scale of Burkina Faso

3.1 Data and method

Regional geological – structural mapping of Birimian terrains by means of aeromagnetic data has proven to be useful to complement field investigations (Vegas *et al.*, 2008; Metelka *et al.*, 2011; Jessell *et al.*, 2012; McFarlane *et al.*, 2019), particularly because the West African context provides uneven lateral outcrop continuity due to an extensive lateritic cover (Chardon *et al.*, 2018). In the present study, we used the aeromagnetic cover of Burkina Faso acquired in the course of the SYSMIN project (System for Mineral Products) at a pixel size of 250 m (see Metelka *et al.*, 2011; Fig. 2a). The image represents the total magnetic intensity, overlaid on a grey-scale shaded relief image of its first vertical derivative (Fig. 2a).

We manually extracted the magnetic form lines (i.e., magnetic fabrics), which are expressed by the ridge-and-trough pattern of the image (Fig. 2a and b). The ridges and the troughs, which can be several tens of km long, mark lateral contrasts in both the total magnetic intensity (color scale) and its derivative (hillshade). At first sight, magnetic fabrics would appear to express a lithological layering of contrasted magnetic properties. However, locally high-angle crosscutting relationships between the foliation and greenstone belt boundaries (Fig. 2b) indicate that it is not only a lithological fabric. Given the generally good agreement between magnetic fabric trajectories and the field measurements of foliation and

cleavage (Fig. 2b), the magnetic fabric is interpreted to mark essentially the foliation or schistosity of the rocks that commonly coincides with the lithological banding to form a composite (“S0-1”) fabric. Extraction of magnetic fabric trajectories on a regional scale by a single operator aims to provide a harmonized view of the strain pattern that would not be biased by an unevenly distributed regolith cover or by the contrasted interpretations of independent cartographers having worked in distant areas. Post-Eburnean mafic dikes have a distinctive signature of very strong, rectilinear and broken anomalies (Fig. 2e). They are omitted on the produced structural map.

The qualitative geometrical analysis of the magnetic fabric traces allowed for the identification of strain gradients indicative of shear zones (Martelat *et al.*, 2000; Chardon *et al.*, 2008). The mapping of those shear zones was made based on the shape of the fabric trajectories and their convergence/divergence patterns, their deflection, truncation, clustering and rectilinearity patterns (Fig. 2b and c). A comparison of the interpretative shear zone traces with those of published maps (Chèvremont *et al.*, 2003; Baratoux *et al.*, 2011) shows an overall good agreement (Fig. 2c and d). But the enhanced lateral continuity of the structural magnetic mapping further allows refining the shear zone pattern by:

- enhancing the kinematic coherency and connectivity of the fabrics and shear zones;
- disregarding earlier interpretative shear zones of too limited length or cutting across continuous and undisturbed magnetic fabric beams to have a kinematic meaning;
- disregarding late- and post-tectonic fractures and faults that have no link to the regional ductile strain field.

Field kinematic criteria are consistent with deflection patterns of the magnetic fabrics along the newly mapped shear zones (e.g., sinistral along the N-trending shear system cutting the entire map area and dextral along the NE-trending shear system in the northeastern part; Fig. 2b and c). The resulting fabric/shear zone map of Burkina Faso (Fig. 3) is complemented in Northern Ghana by the shear zone pattern mapped by Block *et al.* (2016b).

3.2 Fabric and shear zone patterns

The regional structural grain underlined by the fabric trajectories is dominantly NE trending (Fig. 3). The shear zones are mostly arranged as an anastomosed network delimitating lower-strain shear lenses. Almost all shear zones branch onto the neighboring shear zone(s) at an acute angle. This is even the case for the longest shear zones. For instance, the Ouahigouya shear zone bends eastward north of Ouahigouya to be taken into the arched Baraloubé shear zone, which branches onto the NE-trending Gorom Gorom shear zone around 1 W (Fig. 3). The shear zone also merges with the Fada N’Gourma shear zone in Ghana (Block *et al.*, 2016b; see section 4). The systematic anastomosed branching pattern and the connectivity of the shear zones suggest that no individual shear zone did strictly outlast slip along its nearest neighbor(s) and that kinematic coherency has been maintained within the shear zone network throughout the deformation history.

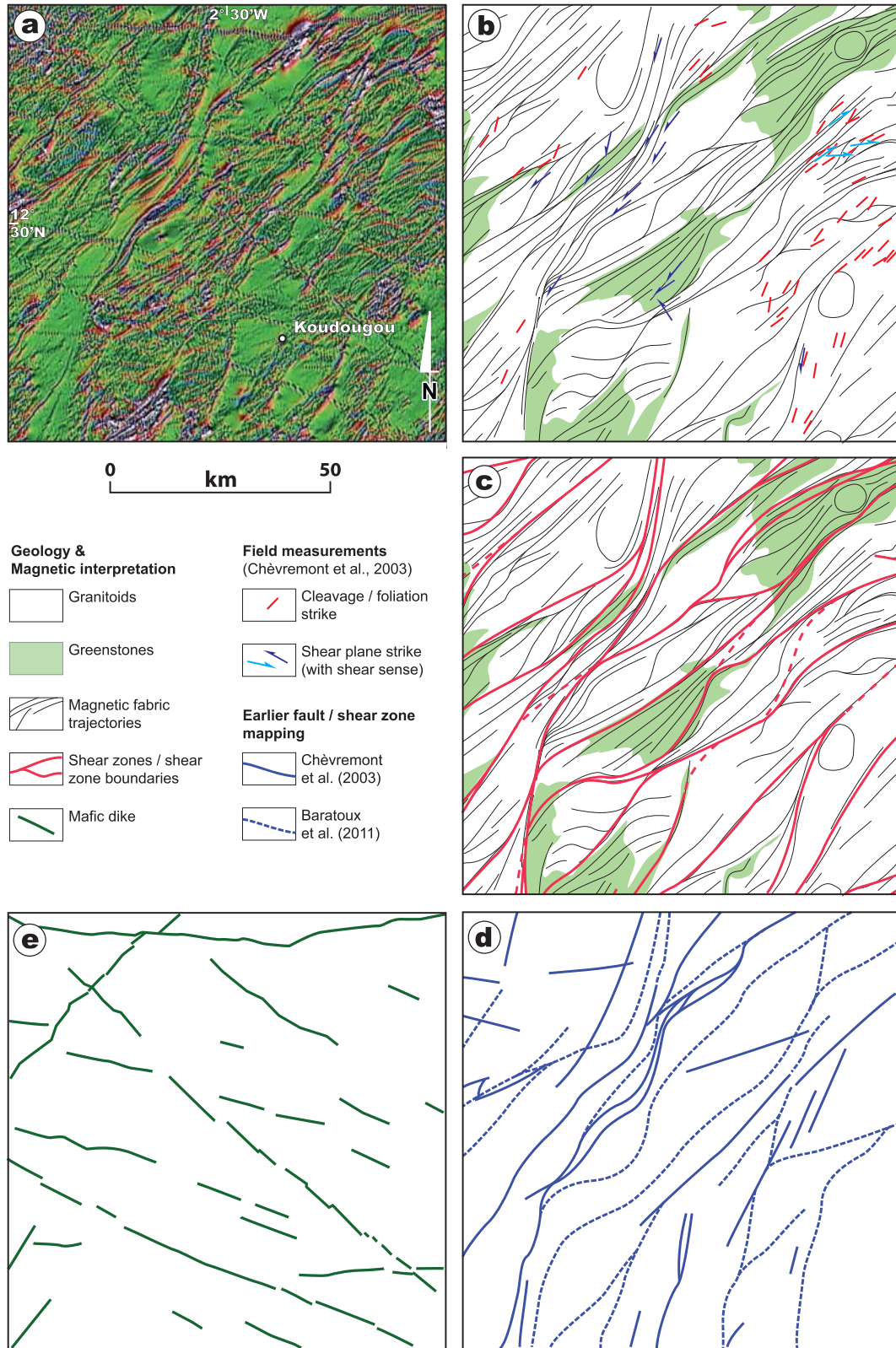


Fig. 2. Fabric/shear zone mapping protocol from aeromagnetic data and its validation illustrated for the Koudougou quadrangle (location on Fig. 1): (a) aeromagnetic map; (b) interpretative fabric trajectories with field measurements of Chèvremont *et al.* (2003); (c) interpretative shear zone pattern. The longest NNE-trending shear zone is the Ouahigouya shear zone (Fig. 3); (d) Faults/shear zones from earlier mapping works (Chèvremont *et al.*, 2003; Baratoux *et al.*, 2011); (e) interpretative pattern of the main mafic dikes. Geological contours are modified from Castaing *et al.* (2003).

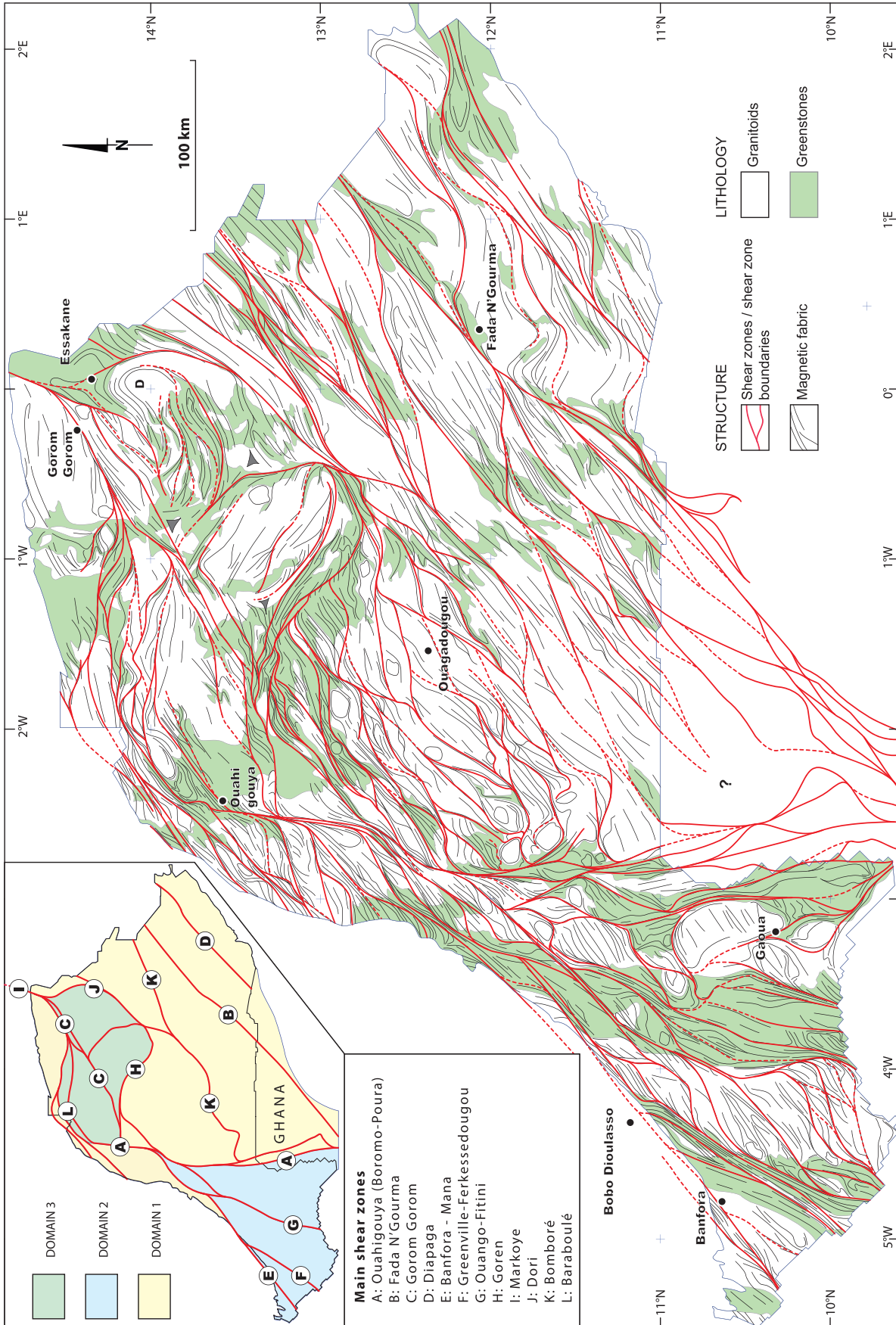


Fig. 3. Fabric/shear zone map of the Birimian basement of Burkina Faso based on aeromagnetic data interpretation. The shear zone pattern in Ghana is adapted from Block *et al.* (2016b). D: Djibo Batholith. The inset shows the structural domains and main shear zones names.

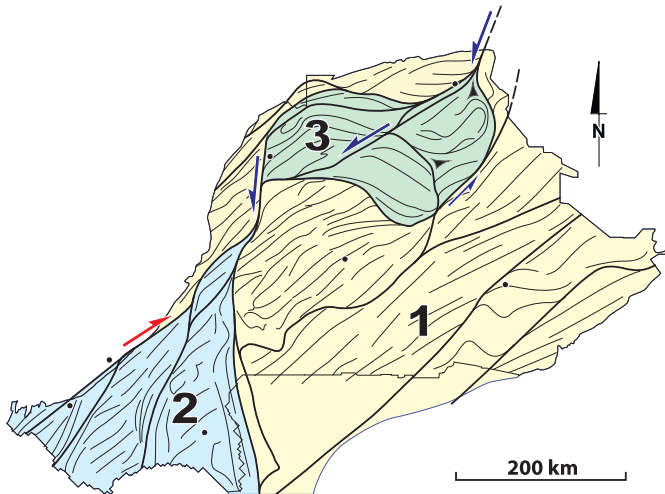


Fig. 4. Sketch structural model of the Eburnean basement of Burkina Faso and Northern Ghana. Colors and numbers refer to the structural domains discussed in the text. Localities of Fig. 3 are shown by dots.

Three structural domains are distinguished at the scale of the country (Fig. 4). Domain 1 occupies half of the exposed Eburnean crust (mostly over the southeastern part of the country; Fig. 4). It is characterized by dominantly rectilinear and parallel NE-trending fabrics and shear lenses that can be up to 200 km-long and 100 km-wide. Deformation was accompanied by the emplacement of batholiths elongate in the direction of the structural grain (Naba *et al.*, 2004). A few straight shear zones of regional length (> 400 km) are identified, the longest one being the Fada N’Gourma shear zone (Figs. 3 and 4). Perturbation of the foliation/shear zone trends South of Fada N’Gourma town is explained by back rotation of the fabrics (Harris, 2003) in between the Fada N’Gourma shear zone and the Diapaga shear zone (Fig. 3). This pattern, originally documented by Vegas *et al.* (2008), attests to a sinistral component of slip along those shear zones. Field observations in mining pits near the Ghana border confirms sinistral kinematics along the Fada N’Gourma shear zone (Woodman *et al.*, 2016). Foliations and cleavages are very steep in both the granitoids and greenstones throughout domain 1 (e.g., Bonnard and Kagambega, 2003; Ouédraogo *et al.*, 2003; Trinquard *et al.*, 2003). Mineral-stretching lineations have variable plunges and may be very steep in the greenstones. Steep lineations in transcurrent shear zones (e.g., Fontaine *et al.*, 2017) are indicative of transpression as the dominant deformation mode of domain 1. If the Fada N’Gourma shear zone appears to be sinistral in Burkina Faso, Block *et al.* (2016b) have shown that it was dextral in northern Ghana (Fig. 3). Sinistral shearing could be restrained to the SSW trending portion of the shear zone between Fada N’Gourma town and the border, as opposed to dextral along its remaining SW trending trace.

According to Block *et al.* (2016b), in situ U-Pb dating indicates that domain 1 crustal shortening and coeval lower crustal flow took place essentially between ca. 2140 and 2110 Ma under high-grade metamorphic conditions, whereas strain localization along the strike-slip shear zones continued afterwards. The regional shear zone located between the Bomboré and Fada N’Gourma shear zones (Fig. 3) functioned

at the time of emplacement of a 2140 ± 7 Ma dioritic body (in situ zircon U-Pb age; Fontaine *et al.*, 2017), indicating that transcurrent tectonics started early in the preserved deformation history of domain 1. Transcurrent tectonics would have lasted until after 2110 Ma (end of high-grade metamorphism) to explain low-grade strike-slip strain localization (< 2100 Ma?).

At the scale of domain 1, fabric trajectories roughly bisect the angle between the two sets of shear zones delimitating the lenses, which is indicative of a component of pure shear achieved by the anastomosed shear zones network (i.e., bulk inhomogeneous pure shear; Gapais *et al.*, 1987). The rectilinearity and parallelism of the fabric trajectories further attest to the intensity and homogeneity of the bulk strain. Shear sense inversion is required along curved lens-bounding shear zones throughout anastomosed shear networks. Kinematic complexities further arise at zipper junctions between shear zones (Hudleston, 1999; Platt and Passchier, 2016) within such networks. Therefore, the protracted convergence period (> 30 My) over which strain field 1 has evolved could result locally in apparently complex kinematic configurations due to shear zones interactions achieving the preservation of the kinematic coherency of the regional structural pattern. Interestingly, the NE-trending shear-fabric pattern could be explained locally on individual shear zones by a wide range of convergence obliquities through time and space, from wrench dominated dextral transpression responding to ENE-trending convergence, to wrench-dominated sinistral transpression (SSE-trending convergence), with intermediate configurations closed to pure shear dominated transpression (SE-trending convergence). In any case, the finite shear-zone/fabric pattern of domain 1 argues for bulk NW-SE coaxial shortening and bulk NE-SW lateral stretching achieved by transpression.

Domain 2 occupies the southwestern and westernmost part of the map area (Fig. 4). It is characterized by an interlinked fabric and shear zone network, which is partitioned between:

- dominant N- to NNE trends within and along the boundaries of the greenstone belts;
- domain 1-type NE trends in intervening granitoid domains.

Overall, domain 2 displays a northward-converging pattern of regional shear zones that end up lining against the Ouahigouya shear zone (Figs. 1 and 3). This shear zone marks the eastern boundary of domain 2 and shows a consistent deflection pattern of foliation north of 12° N, attesting to a component of finite sinistral slip (Fig. 3). Fabric deflection patterns along the NNE-trending shear zones are ambiguous, with dextrally asymmetrical trajectories, symmetrical trajectories (i.e., pure shear?) or –less commonly– sinistrally asymmetrical trajectories, sometimes along strike a same shear zone. NE-trending shear zones tend to show more consistently dextral asymmetrical deflection patterns. The density of those dextral NE-trending shear zones appears to increase north-westward up to the longest dextral shear zone of the domain (i.e., Banfora shear zone; Fig. 3). Baratoux *et al.* (2011) interpreted the regional shear zone pattern in domain 2 as due to a D2, wrench-dominated transpressional deformation event, which followed E-W shortening having led to the steep longitudinal foliation pattern of the greenstone belts. But the steep fabrics of greenstone belts must have functioned together

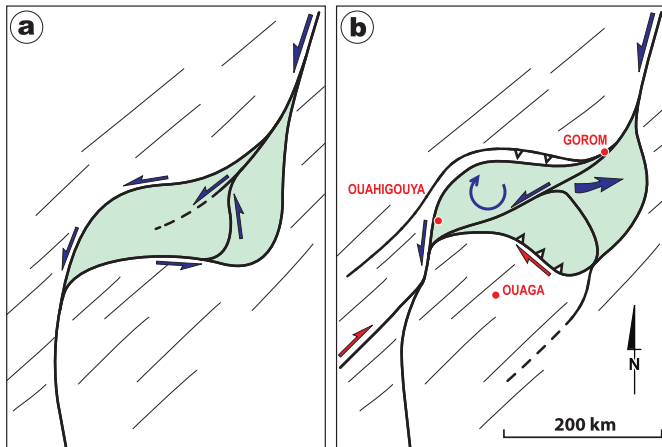


Fig. 5. Kinematic model for the growth of the Bouroum-Goren restraining bend in between the Markoye shear zone (to the East) and the Ouahigouya shear zones (to the west): (a) early bending stage; (b) final configuration. The model implies the closure of zipper junctions (Platt and Passchier, 2016) at both tips of the restraining bend.

with the shear zones and the granitoid fabrics for a protracted period to maintain the kinematic coherency of the anastomosed shear pattern between N-S greenstone corridors and intervening granitoids (Fig. 3). As documented by Lompo (2010), lineation pitches are variable on greenstone foliations (attesting to flattening strain) and are preferentially high (attesting to vertical stretch). Steep lineations are considered by Lompo (2010) to result from vertical tectonism due to diapirism. However, kinematic criteria near the granite-greenstone contacts that would substantiate such a model (Chardon *et al.*, 1996) are still missing. The greenstone foliation and lineation pattern would instead be indicative of pure shear dominated transpression (Tikoff and Greene, 1997; Dewey *et al.*, 1998). The fact that granitoids show preferentially shallowly plunging lineations outside pluton aureoles (Lompo, 2010; Baratoux *et al.*, 2011) argues for strain partitioning in the anastomosed shear network between pure shear dominated transpression in the greenstone belts and strike-slip dominated transpression in the granitoids. U-Pb geochronology indicates that the domain 2 strain pattern was functional after ca. 2160 Ma (oldest TTG and plutons interpreted to have emplaced during acquisition of the steep N-trending fabric) and until after ca. 2100 Ma (latest potassic plutons affected by mylonitic shearing; Baratoux *et al.*, 2011). There is therefore a significant time overlap (> 40 My?) in the deformation history of domains 1 and 2.

Domain 3 occupies the north-central part of the country (Fig. 4) and corresponds to an area of domain 1 where the regional NE structural trend of the fabrics, shear zones and greenstone belts is perturbed. The domain is a composite, lazy Z shaped mega shear lens. It is divided in two by the Gorom Gorom shear zone, which is also the symmetry axis of the mega shear lens and connects the sinistral Markoye shear zone (Tshibubudze and Hein, 2013), in the NE, to the northern sinistral portion of the Ouahigouya shear zone, in the SW (Fig. 3). The mega shear lens is interpreted as a restraining bend developed in between the two shear zones (Fig. 5) and is called hereafter the Bouroum-Goren restraining bend.

This interpretation accounts for the map scale bend made by the Goren greenstone belt as the result of clockwise rotation concomitant with sinistral slip of the northward propagating tip of the Ouahigouya shear zone (Fig. 5). Rotation and slip produced the abutment pattern of domain 1-related NE-trending foliations against NE trends of the southwestern envelop of the restraining bend NNE of Ouagadougou (against the Goren shear zone; Figs. 3 and 5). This abutment pattern earlier documented by Hein *et al.* (2004) is here interpreted to result from dominantly dextral transpression along the Goren shear zone as a consequence of growth and rotation of the restraining bend (Fig. 5). Dextral shearing combined with NE-side up reverse slip documented by Hein (2010) along that shear zone is therefore consistent with the restraining bend model (Fig. 5). The model further accounts for NNE-directed thrusting combined with sinistral strike-slip documented by Nikiema (1992) and Debat *et al.* (2003) along the northern margin of the restraining bend (Baraboulé shear zone; Fig. 3) in a context symmetrical to that of the Goren shear zone (Figs. 3 and 5). Reverse shearing along that margin post-dates the amphibolite facies metamorphic aureole of the plutons (Debat *et al.*, 2003). This, together with the abutment pattern at the southwestern margin of the bend, shows that the Bouroum-Goren bend is a late tectonic feature that has grown at the expense of domain 1.

3.3 Kinematic interpretation

The Eburnean deformation pattern of Burkina Faso (Figs. 3 and 4) is interpreted to result from the interferences between the strain field typified by domain 1 and regional-scale shear zones trending N to NNE. Strain field 1 developed throughout the country. It was achieved by steep and parallel foliations trending NE combined with a regionally anastomosed network of transpressive shear zones, which recorded bulk inhomogeneous shortening on a regional scale. The first type of perturbation in that strain field occurs in domain 2, where N- to NNE-trending shear systems alternate with corridors of NE-trending, domain 1-type fabrics (Fig. 4). This interference results from the activation of the northward-convergent fan pattern of regional-scale shear zones (i.e., Fig. 1) onto strain field 1. The second type of perturbation was produced by the development of the Bouroum-Goren restraining bend between the northern tip of the regional-scale fan shear zone pattern of domain 2 and the Markoye shear zone (Figs. 3 and 5). Domain 1 would therefore contain the oldest Eburnean pervasive fabrics preserved regionally. However, perturbation of strain field 1 instead of its regionally pervasive overprint by N- to NNE trending domain 2-type shears argues for the rejuvenation of the NE-trending structural grain during the development of patterns 2 and 3 despite potential change(s) in the tectonic boundary conditions the later might have been triggered by. This could further explain local kinematic complexities expected inside domain 1.

4 Craton-scale deformation pattern

4.1 Fabric and shear zone patterns

We produced a new fabric/shear zone map of the sWAC based on fabric trajectories derived from foliation/cleavage

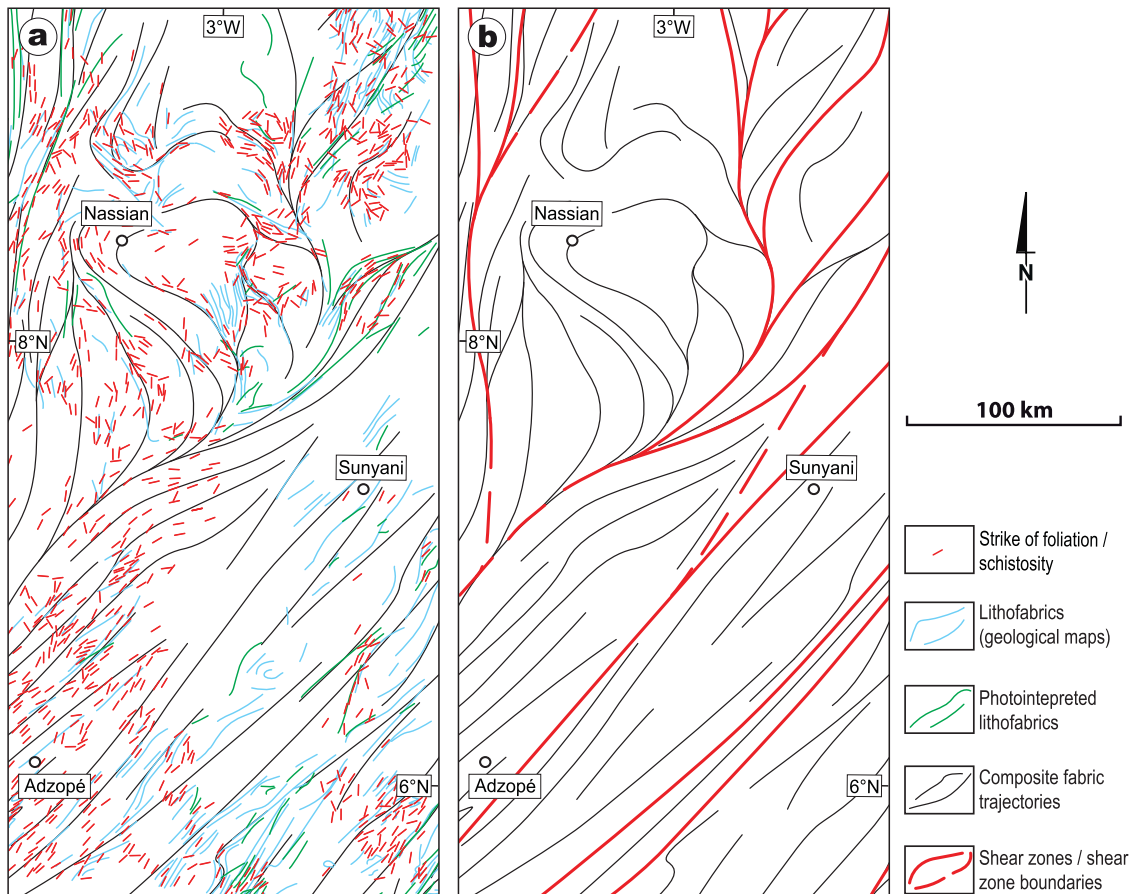


Fig. 6. Fabric/shear zones mapping protocol at the scale of the southern West African craton (Fig. 7a) illustrated on the Comoé area in the Côte d'Ivoire – Ghana border region (location on Fig. 1): (a) fabric trajectories drawn after foliation measurements, lithological fabrics from geological maps and photointerpreted fabrics; (b) interpretative shear zone pattern superimposed on the fabric trajectories. Source maps: Bureau de Recherches Géologiques et Minières (1963), Eisenlohr (1989), Delor *et al.* (1992, 1995), Adou *et al.* (1995), Siméon *et al.* (1995), Zeade *et al.* (1995), Griffis Consulting and Exploration Services, Watson Geophysics (1998), Zitzmann (1998), Ghana Geological Survey (2008a, 2008b, 2008c).

data and lithological fabrics retrieved from geological maps and photointerpretation using the GoogleEarth™ software (Figs. 6 and 7a). The main shear zones were drawn from the fabric trajectories the same way shear zone were derived from magnetic fabrics in the case of Burkina Faso (Figs. 2 and 6).

Based on the craton-scale shear zone patterns (Fig. 7a), five deformation provinces are defined (Fig. 7b). The Burghana province has consistent NE-trending transpressive fabrics and shear zones (Fig. 7b). The Ivorian province is defined by a lozenge-shaped shear zone pattern covering most of Côte d'Ivoire (Fig. 7b). It consists in stacked, N- to NNE- trending shear lenses that end up in the Bouroum-Goren restraining bend in Burkina Faso (domains 2 and 3 in Figs. 4 and 7b). The Archean province is affected by a highly curved, eastward-convex shear zone pattern (Fig. 7b). The southern flank of that arch is made of dominantly dextral shears and the northern flank by a sinistral shear system (Egal *et al.*, 2002; Fig. 7b). The hybrid province (Fig. 7b) has undergone mostly Eburnean high-grade metamorphism (Pitra *et al.*, 2010) and consists of reworked Archean crust and hybrid Archean-Birimian crust (Kouamelan *et al.*, 1997, 2018). It is affected by ENE-trending shear lenses pinched south of the Archean province against the

Ivorian province (Fig. 7b). Two shear lenses of that province (mostly Birimian) may be mapped north of the contact between the Archean and Ivorian provinces (Caby *et al.*, 2000; Thiéblemont, 2016; Fig. 7b). The Guineo-Malian province corresponds to a northward fanning shear zone pattern in between the Archean and Ivorian provinces (Fig. 7b).

4.2 Kinematic interpretation

The early NE-trending regional fabrics are inferred to have affected at least the Ivorian province in addition of the Burghana province (K. Traoré, unpublished data; see also Fig. 4). Development of the Ivorian shear pattern was due to the northeastward movement of the Kénéma-Man domain and dextral slip along a wide NE-trending deformation belt belonging to the Burghana province (Figs. 7 and 8). This belt is called hereafter the Burghana-Transamazonian (B-T) mega shear because it has to extend southwestward into the Guiana shield where the South American counterpart of the Eburnean orogeny is called Transamazonian (e.g., Ledru *et al.*, 1994; Vanderhaeghe *et al.*, 1998). The B-T mega shear marks the southeastern contact of the orogen with an Archean buttress

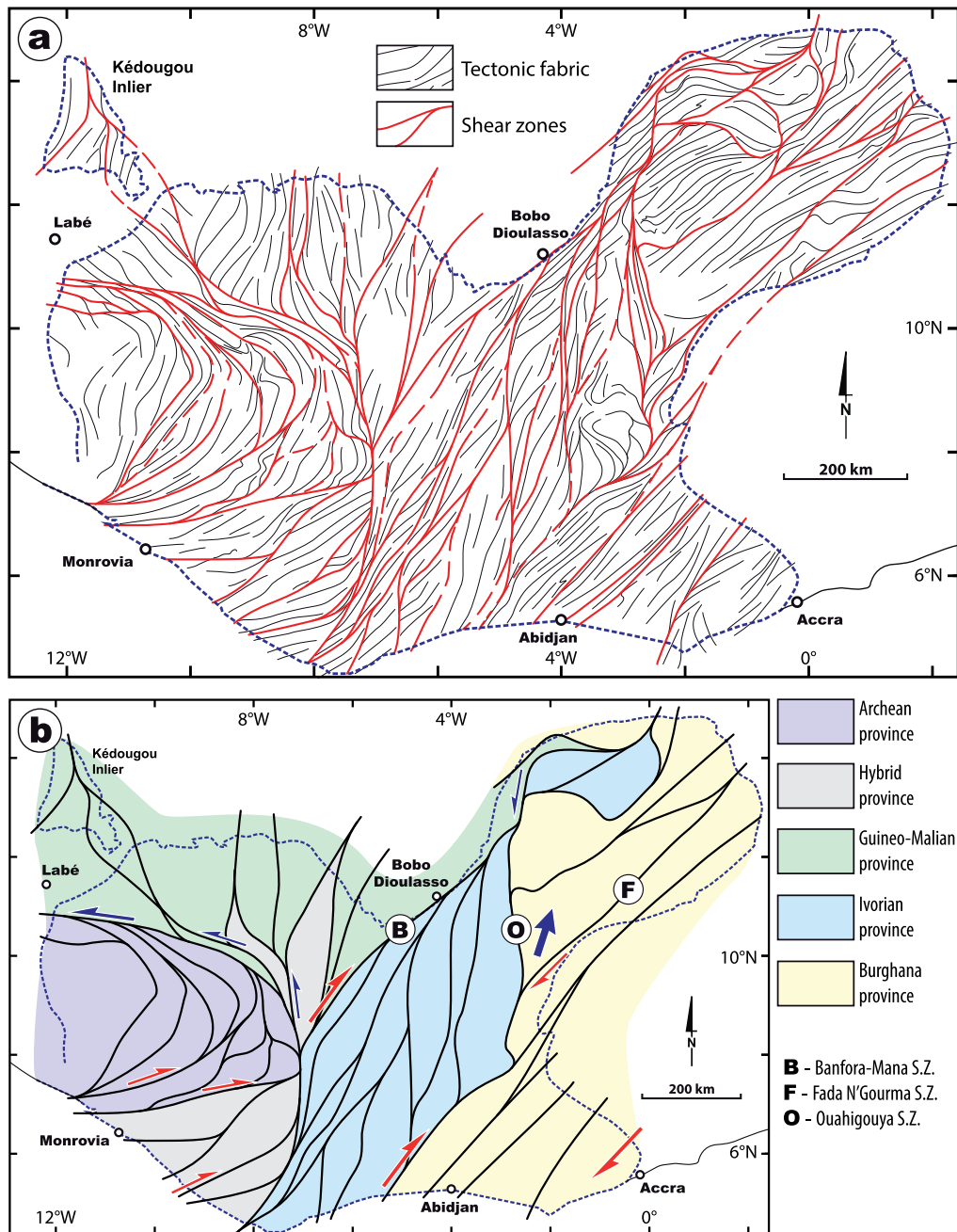


Fig. 7. (a) Fabric/shear zone map of the southern West African craton after a compilation of fabric data from published geological maps and photointerpretation (Fig. 6). This new map allows refining and complementing the shear zone/fault pattern of Milési *et al.* (1989) and Thiéblemont (2016) (see Fig. 1 for comparison); (b) deformation provinces map of the southern West African craton based on (a). The boundary between the Guineo-Malian and the Burghana provinces in northwestern Burkina Faso (NE of Bobo Dioulasso) is not well-constrained.

(Feybesse and Milesi, 1994; Ledru *et al.*, 1994; Fig. 8) that could be the Congo craton at the time (e.g., Lompo, 2009). Structural studies of the B-T mega shear attest to a protracted and locally complex deformation history (Perrouy *et al.*, 2012; McFarlane *et al.*, 2019 and references therein). But the systematic interference between steep reverse and strike-slip shearing (Ledru *et al.*, 1994; Feybesse *et al.*, 2006) argues for deformation partitioning through time and space during long-lasting or repeated transpression (Eisenlohr and Hirdes, 1992)

comparable to domain 1 of Burkina Faso (Figs. 1 and 4). Strain analysis of a 80-km wide strip of the northwestern B-T mega shear suggests > 400 km of dextral or simple shear (Jessell *et al.*, 2012), which is consistent with the deflection pattern of fabrics and shear zones of the Ivorian province towards that contact and the net dextral slip suggested by our reconstruction (Figs. 6–Figs. 6, 7a and 8).

Collision with the Kénéma-Man domain drove the growth of the N-S strike slip duplexes of the Ivorian

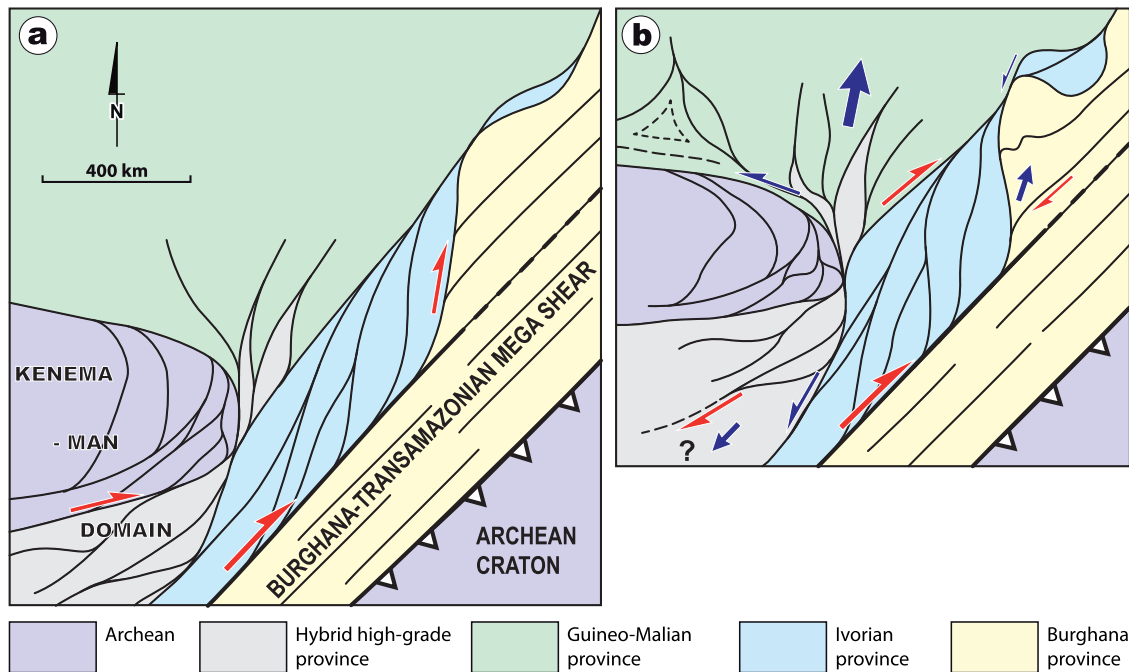


Fig. 8. Kinematic model for the late development stages of the Eburnean orogen of the southern West African craton. The Burghana-Transamazonian mega shear and the adjoining Archean craton have been taken as stationary in this reconstruction: (a) early shortening of the orogen by the Kénéma-Man domain against the Burghana-Transamazonian mega shear: nucleation of the Ivorian lozenge shear zone pattern; (b) final configuration: shortening of the Ivorian shear lozenge and northward extrusion of the Guineo-Malian domain; southwestward extrusion (?) of the high-grade hybrid crust of the Kénéma-Man domain.

province while the Burghana province maintained its NE-trending structural grain (e.g., domain 1 versus domain 2 in Burkina Faso; Fig. 4). Sinistral kinematics of the N-trending portions of the Ivorian province shear zones (e.g., Mortimer, 1992; Vidal *et al.*, 1996; Gasquet *et al.*, 2003; Grenholm *et al.*, 2019) accommodated conjugate slip with respect to the B-T mega shear and the Banfora-Mana shear zone (Fig. 7b), contributing to shortening and northeastward extrusion of the province (Fig. 8). That extrusion was ultimately absorbed in Burkina Faso by escape in between the Fada N’Gourma and Ouahigouya shear zones (see also Grenholm *et al.*, 2019) and formation of the sinistral Bouroum-Goren restraining bend at the tip of the Ivorian lozenge (Figs. 7 and 8).

Over-bending of the arched shear pattern in the Archean province is suggested during collision with the Birimian crust (Fig. 8), resulting in an apparent orocline. Bending was associated with transpressional crustal thickening and high-grade metamorphism (Pitra *et al.*, 2010) in the southeastern Kénéma-Man domain, while voluminous magmatism and sinistral transpression took place along the northern flank of the bend (Égal *et al.*, 2002; Eglinger *et al.*, 2017). Exhumation of high-grade terrains in the Archean and hybrid provinces is likely linked to transpressional thickening and/or lateral escape on both sides of the contact zone between the Archean and Ivorian provinces (Fig. 8). The Guineo-Malian province developed a shear zone pattern mostly indicative of northward extrusion from the collision zone between the Archean and Ivorian provinces (e.g., Feybesse and Milesi, 1994; Fig. 8). The kinematic reconstruction implies that significant N-S convergence (> 200 km) must be absorbed at the northern margin of the Kénéma-Man domain as a consequence of slip

along the B-T mega shear (Fig. 8). This could have required subduction to have taken place along this boundary to absorb that convergence (see Eglinger *et al.*, 2017 and Wane *et al.*, 2018 for recent petro-geodynamical scenarios).

The deformation of each Birimian province was mostly controlled by its position in the orogen. The B-T mega shear remained a rectilinear boundary of the orogen. The rest of the Burghana province maintained its dominant NE-trending Eoeburnean tectonic grain during collision with the Kénéma-Man domain because it was the most distant from the collision zone (Figs. 7 and 8). A shift in peak age felsic plutonic activity, from older to younger, is documented across the Ouahigouya shear zone (Block *et al.*, 2016a) and the Banfora-Mana shear zone (Parra-Avila *et al.*, 2017) (Fig. 7b). This also suggests stiffness contrasts among the Burghana (outside the B-T mega shear), Ivorian and Guineo-Malian provinces that could also explain their contrasted mechanical/kinematic behavior during collision. The sWAC finite strain pattern was acquired mostly from ca. 2140 Ma onward. Regional deformation, strain localization on shear zones and plutonism lasted until ca. 2060 Ma with an overall younging age in plutonism towards the collisional contact with the Kénéma-Man domain (Grenholm *et al.*, 2019).

5 Significance of large-scale shear zone patterns: a discussion

Although the regional and craton-scale strain and shear zone patterns of the sWAC evolved over tens of millions of

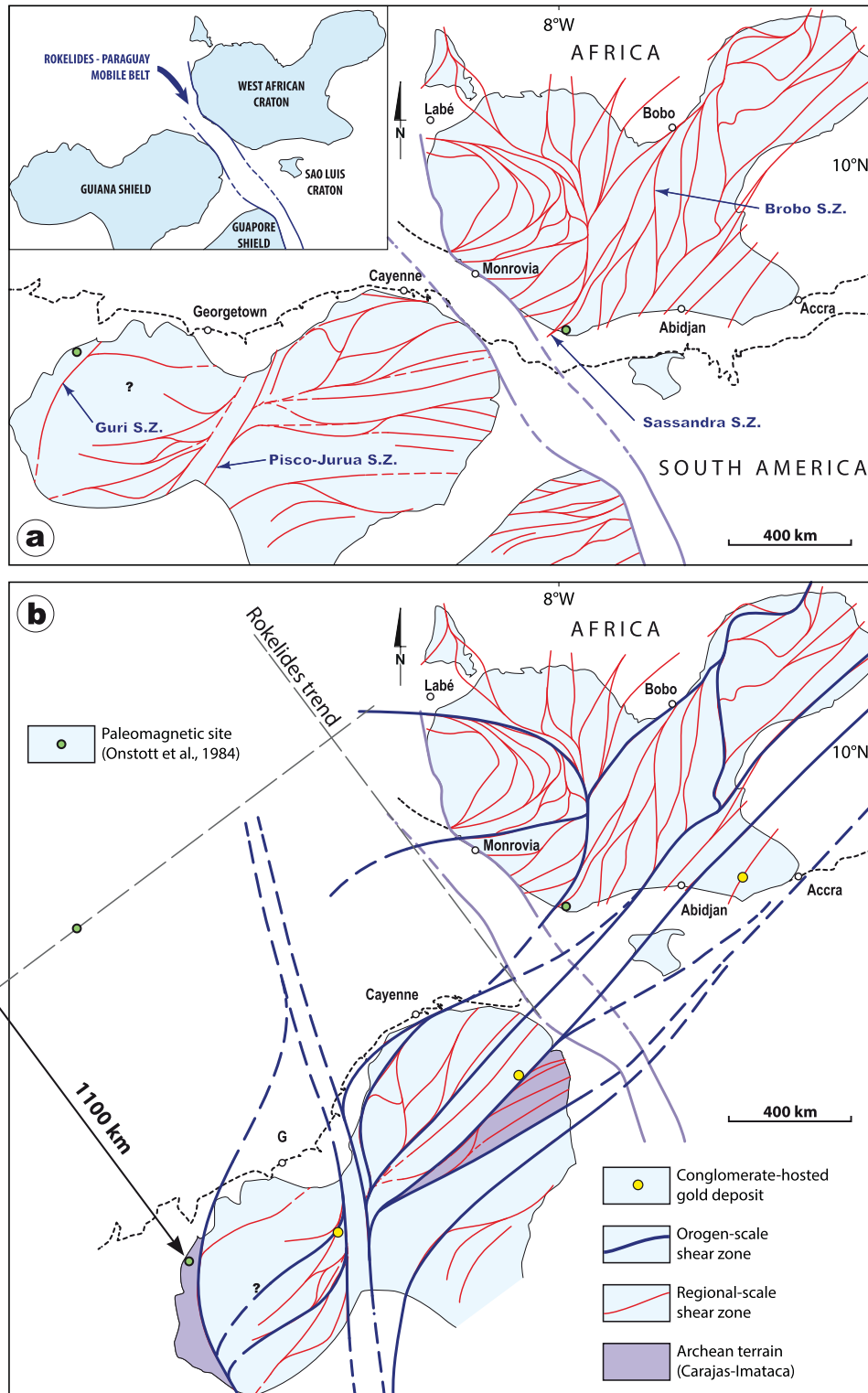


Fig. 9. Transatlantic Eburnean-Transamazonian correlation (in today’s Africa reference frame): (a) configuration at the end of the Paleozoic based on the Africa – South America pre-rift fit of the Equatorial Atlantic Ocean of [Ye *et al.* \(2017\)](#). The shear zone pattern in South America has been compiled mostly after [Gibbs and Baron \(1993\)](#), [Delor *et al.* \(2003b\)](#) and [Cordani *et al.* \(2006\)](#). The inset shows the relative position of cratons and the Panafrican mobile belt; (b) reconstruction of the Eburnean-Transamazonian orogen before Neoproterozoic formation of the Rokelides – Paraguay mobile belt (see text for further explanation). The reconstruction implies chisel-like closure of a Panafrican ocean accompanying clockwise rotation of the Guiana shield and limited (< 300 km) right-lateral movement along the mobile belt during Panafrican collision. Chisel-like closure of a “Rokelides ocean” would explain why a suture zone is imaged along the NNE-trending segment of the orogen (South of Labé; [Ponsard *et al.*, 1988](#)), whereas such a suture is disputed further to the SE (Monrovia area; [Williams and Culver, 1988](#)).

years, their geometrical interpretation provides first-order kinematic constraints on the evolution of the orogenic system. The sWAC is imprinted by bulk inhomogeneous transpressional shortening recorded by anastomosed networks of steep shear zones and steep regional planar fabrics. Anastomosed shears define shear lenses up to ca. 150 and 300 km in length on the regional and craton-scale, respectively. The connectivity of the shear networks implies limited horizontal and vertical offsets of individual shears. The map view conjugate and mostly symmetrical character of the shear lenses show that lateral horizontal stretching with a coaxial component of deformation is significant. Such regional strain patterns are indicative of the homogeneous shortening of a soft orogenic lithosphere lacking a stiff mantle layer, which would be typical of Precambrian accretionary orogens (Choukroune *et al.*, 1995; Gapais *et al.*, 2005; Cagnard *et al.*, 2006a, 2006b; Cruden *et al.*, 2006; Chardon *et al.*, 2009; Harris and Bédard, 2014). Instead of large localized displacements, shear zone networks accommodated bulk deformation and absorbed kinematic incoherencies generated by instabilities or heterogeneities in the crustal flow pattern. Only a few very long (> 300 km) rectilinear shear zones such as the B-T mega shear may have accommodated displacement of prime tectonic importance between deformation provinces (Fig. 8).

5.1 Refining the shear zone patterns: kinematic consequences

Mapping of tectonic fabrics has allowed refining the regional shear zone patterns. The kinematic information retrieved from those patterns may be overlooked by conventional geological mapping or local structural geology, as illustrated on the Burkina Faso case example. Our map (Figs. 3 and 7) indicates that the branches of the Banfora-Mana shear system (Fig. 7b) do not terminate in central Burkina Faso as shown on Figure 1 but merge and branch onto the Ouahigouya shear zone (see Castaing *et al.* (2003), Figs. 3 and 7b). Instead of being the latest (D3) tectonic features in the sWAC as argued originally by Milési *et al.* (1989), those dextral shear zones have coexisted with N-S, dominantly sinistral conjugates in the interlinked Ivorian shear lozenge (Figs. 1 and 7). Those dextral shears could not therefore have rotated the Goren greenstone belt and shear zone (Fig. 3) from N-S to NW-SE as suggested by Feybesse *et al.* (1990). Hein (2010) interpreted dextral transpression in the Goren shear zone (Figs. 3 and 5) as due to an Eoeburnean (i.e., “Tangaeen”) regional phase of SW-vergent reverse faulting (Tshibubudze *et al.*, 2009). However, the abutment pattern of domain I fabrics against the Goren shear zone (Figs. 3–5) indicates that domain I NE-trending fabrics had to preexist the shear zone (and acquisition of the NW Goren trend). The Goren shear zone cannot therefore be an Eoeburnean feature. The sinistral restraining bend model therefore appears to better account for the regional finite strain pattern and field kinematic observations made in Central Burkina Faso. For Feybesse *et al.* (1990), steep tight to isoclinal folds in greenstone belts were originally produced by Eoeburnean low-angle thrust tectonics. Still, if earliest pervasive fabrics of regional extent in Burkina Faso should preferentially be preserved in domain I (Fig. 4), evidence for an even earlier (Eoeburnean)

kinematically distinct regional deformation – if any – would be elusive given the imprint of long-lasting transpression along the NE structural grain.

5.2 The deformation phase approach

The present work suggests that the deformation phase approach that would not take into account large-scale (100–1000 km) structural patterns could be misleading, as apparently complex kinematic configurations arise from interactions among shear zones (e.g., Fig. 5) that would classically be systematized as a series of deformation events. But more generally, the phase approach would be misleading for four reasons. The first reason is that even a single progressive simple shearing episode can generate the superimposition of apparently incompatible classes of schistosity, veins and folds that rotate differentially during deformation (Brun and Choukroune, 1981) as well as conjugate shear zone patterns (Harris and Cobbold, 1985). Simple shearing also generates C' shears oblique on the fabrics tracking the bulk shear direction (Berthé *et al.*, 1979). In some cases, the phase approach would lead to interpret outcrop- or map-scale C' shears as resulting from a deformation event younger than, and kinematically unrelated to, the shear deformation that actually produced them. C' shears, as syn-cooling features, are indeed commonly brittle-ductile and contrast with the simple shearing fabrics they affect. Such a coexistence of brittle and ductile structures during progressive deformation tends to be ignored in the phase approach. The second reason to challenge the phase structural approach is the ubiquitous three-dimensional strain partitioning between strike-slip shearing, reverse shearing and planar and linear fabrics that coexist in transpressional contexts (Hudleston *et al.*, 1988; Tikoff and Teyssier, 1994; Gapais *et al.*, 2005, 2018; Harris and Bédard, 2014). The third reason is that by favoring a local phase approach, one tends not to consider the spatial kinematic diversity of localities as part of a progressive/bulk deformation pattern that would be coherent on a regional scale (examples of regional kinematic integration are provided, for instance, by Martelat *et al.*, 2000; Gapais *et al.*, 2005; Harris and Bédard, 2014; Cochelin *et al.*, 2017). The last and fourth reason is that one must consider fabrics and shear zones rotation and the zipping/unzipping of shear zones (Platt and Passchier, 2016) during the deformation history of regional structures (e.g., Bouroum-Goren bend; Fig. 5). Those processes would commonly lead to shear sense inversion along shear zones through time and/or to spatially contradictory shear senses along individual shear zones at a given time, notwithstanding the fact that the pure shear component of transpression tends to generate conjugate strike-slip shears in individual shear zones.

Dynamic analysis, on which the phase approach is implicitly based, would therefore be of limited value (Tikoff and Wojtal, 1999; Gapais *et al.*, 2000) to infer shortening directions within shear systems such as that of the sWAC (see Grenholm *et al.*, 2019 for a compilation of such directions). The concept of episodic Eburnean orogeny (Block *et al.*, 2016b; McFarlane *et al.*, 2019) derived from such analyses relies on unnecessarily complex and untestable deformation histories, which are based on 5/7-phase scenarios ignoring partitioning of protracted convergent deformation that led to the finite strain pattern. Only

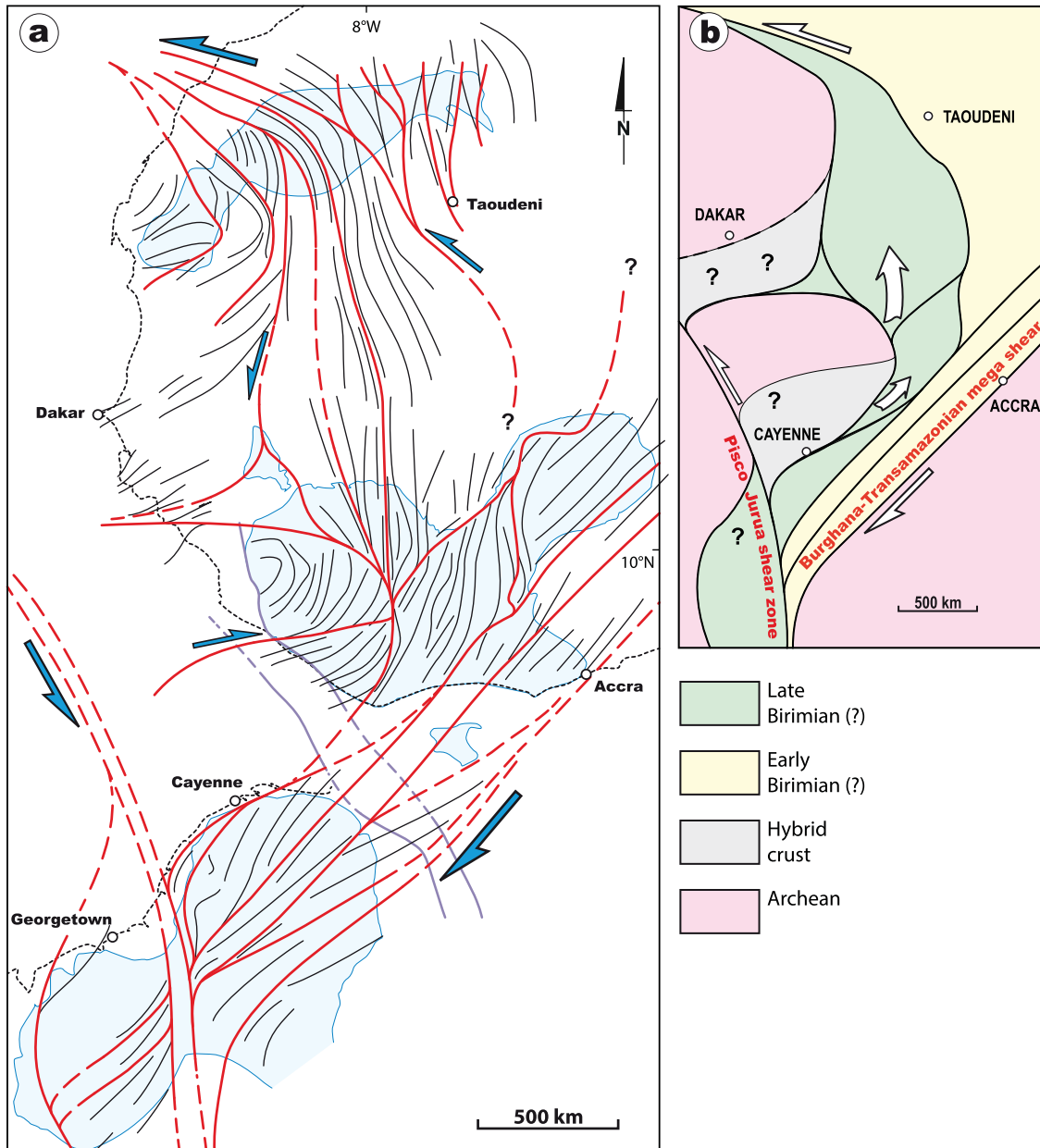


Fig. 10. (a) Structural pattern of the Eburnean-Transamazonian orogen. Fabrics are in black and the shear zones in red. Fabrics in Africa (adapted from Jessell *et al.*, 2016) derive from joint interpretation of medium-wavelength (> 30 km) Bouger gravity anomaly and magnetic data in exposed and buried Archean and Birimian crust. Shear zones in the Reguibat shield (northern outcrop of the West African craton) are adapted from Milési *et al.* (2010); (b) corresponding tectonic model (same frame as Fig. 10a).

structural artifacts that would definitely not fit into the finite strain pattern could provide clues on an early tectonic history but the distortion and rotation they must have undergone would be difficult to take into account.

5.3 Tectonic significance of the shear zone patterns

In map view (e.g., Fig. 7a), fabrics and shear zones track the horizontal component of finite elongation and the lines of no finite elongation, respectively (Gapais *et al.*, 1987, 1991).

Thus, they have no simple relationships to the degree of slip partitioning or shear sense along shear zones or to a far-field convergence vector. NW-trending shortening inferred from local data over the Ivorian and Burghana provinces (Grenholm *et al.*, 2019; Fig. 7b) accounts for the fairly consistent NE trend of the fabrics (Fig. 7a). This fabric trend actually records the pure shear component of transpression parallel to the B-T mega shear bounding the orogen (e.g., Fig. 10a), not actual convergence with the southeastern craton. Relative displacement between the Kénéma-Man domain and the southeastern craton would be E- or NE-trending (e.g., Fig. 8), instead of

NW-trending as deduced as normal to the NE-trending fabrics (Grenholm *et al.*, 2019). Therefore, strain partitioning inside and at the margins of the orogen prevents retrieval of convergent plate kinematics and slip sense on a shear zone has no simple relationship to the relative convergence vector. In other words, the shear zone patterns are indicative of the flow regimes inside the orogen, which are primarily imposed by the shape of the stiff converging provinces to the soft Birimian lithosphere.

5.4 Eburnean orogenic processes

Syn-convergence extensional exhumation of partially molten lower crust (e.g., Block *et al.*, 2015, 2016; McFarlane *et al.*, 2019) took place at a low angle to the regional structural trends in the sWAC, attesting to the involvement of the lower crust in the lateral flow process (e.g., Chardon *et al.*, 2011). Exhumation was also permitted despite convergence in sites kinematically favorable for transensional exhumation to take place (e.g., Cochelin *et al.*, 2017) such as the junction of regional shear zones (e.g., between the Ouahigouya and Fada N'Gourma shear zones; Block *et al.*, 2016b; Fig. 7b). Lateral flow is interpreted to have compensated part of the crustal thickening of the soft Birimian crust.

Whether early Eoeburnean deformation could proceed by low-angle thrust stacking of mafic crustal slices and supracrustals or already by steep transpression remains an open question. The first tectonic process could explain early HP-LT apparent geotherms recorded by greenstones (Ganne *et al.*, 2012, 2014; Block *et al.*, 2015). But transpressional burial of supracrustals by pop-down tectonics (Gapais *et al.*, 2014; Gapais, 2018) and subsequent upward extrusion (Thompson *et al.*, 1997) might as well and would be consistent with early and long-lasting transpression leading to the sWAC finite structural pattern. If PTT paths of Birimian supracrustals may be “reminiscent of modern collisional orogens” (McFarlane *et al.*, 2019), the pervasive transpressive structural pattern of the Eburnean orogen does not compare at all the thrust or gravitational collapse patterns of Variscan or Alpine collision belts.

6 Configuration of the Eburnean-Transamazonian orogen

The updated shear zone map of the sWAC (Fig. 7a) prompted a new reconstruction of the Eburnean-Transamazonian orogen between the sWAC and the Guiana shield across the Equatorial Atlantic Ocean (Fig. 9). Existing reconstructions (e.g., Feybesse and Milesi, 1994; Ledru *et al.*, 1994) were based on a pre-rift fit of Africa and South America. They did not consider the Neoproterozoic relative displacement between the two cratons (Fig. 9a) suggested by paleomagnetic data. Onstott and Hargraves (1981) and Onstott *et al.* (1984) indeed computed ca. 1000 km of sinistral translation of the cratons parallel to the Panafrican Rokelides-Paraguay mobile belt to allow for both aligning the Guri and Sassandra shear zones and merging the discordant paleopoles from the two Birimian localities shown on Figure 9a. Later reconstruction of Nomade *et al.* (2003) connected the Guri shear zone to the Sassandra shear zone and the Pisco-Jurua shear zones to the Brobo shear zone (Fig. 9a). But this reconstruction generated

incoherencies in the transatlantic Eburnean strain pattern (Nomade *et al.*, 2003) and was incompatible with the geometry of the Panafrican mobile belt. Indeed, the Guiana shield would be transected in the middle by the mobile belt in this reconstruction, whereas it must have lied southwest of the belt before the Neoproterozoic. Furthermore, the Brobo shear zone, which is cut by the B-T mega shear (Figs. 8 and 9), could not extend southward into the Pisco-Jurua shear zone and its flanking granulite belt that form the latest (< 2.05 Ga) structure cutting the entire Guiana shield (Delor *et al.*, 2003a; Kroonenberg *et al.*, 2016).

The new reconstruction (Fig. 9b) is based on a recent pre-rift continental fit taking into account syn-rift extensional deformation of the continental margins (Ye *et al.*, 2017) and a structural correlation that allows aligning the longest and most rectilinear shear corridor in the Guiana shield with the B-T mega shear. This shear corridor is flanked to the SE by (or hosts) an Archean Carajas-Imataca unit (e.g., Delor *et al.*, 2003b; Cordani *et al.*, 2006; Fig. 10b), in agreement with the B-T mega shear having interacted with an Archean craton located to the SE (Fig. 8). Our reconstruction would also be consistent with the B-T mega shear being the sole host of conglomerate-hosted gold deposits (Tarkwa-type; e.g., Frimmel, 2014) in the orogen (Fig. 9b). The restoration requires 37° anticlockwise rotation of the Guiana shield, 270 km of net sinistral slip along the mobile belt, and implies 1080 km of sinistral offset of the paleomagnetic sites of Onstott and Hargraves (1981) parallel to the mobile belt (Fig. 9b). The proposed reconstruction therefore accounts for:

- an updated structural-geological Eburnean-Transamazonian correlation;
- paleoproterozoic paleomagnetic data;
- an updated Equatorial pre-rift fit of Africa and South America.

The new reconstruction was used to compile a structural map at the scale of the West African craton and the Guiana shield (Fig. 10a) on the basis of the present work and the geophysical data compiled and interpreted by Jessell *et al.* (2016). This map shows that the western portion of the Eburnean-Transamazonian orogen is arched as a result of indentation by two Archean provinces (Fig. 10a). The orogen is pinched in the South in between the B-T mega shear and the Pisco-Jurua shear zone, the later having acted as a late structure necking the orogen (Figs. 9 and 10). Symmetrical pinching is suggested at the northern termination of the arch although later than in the sWAC, as suggested by geochronological data (Grenholm *et al.*, 2019). The new orogen-scale model would explain northward escape/extrusion of the Guineo-Malian crust into the central part of the orogenic arch where it would be the widest (Figs. 8 and 10b). Continuation of the orogen south of the map area of Figure 10 as an Eburnean megaboudin is not precluded.

7 Conclusions

Regional-scale (100–1000 km-long) shear zones of the Eburnean accretionary orogen are arranged as a transpressional

anastomosed network that accommodated long-lasting distributed shortening and lateral flow of the hot Birimian lithosphere between the converging Kénéma-Man and Congo Archean provinces. Interference patterns in the shear network were due to interactions among shear zones and three-dimensional partitioning of progressive transpression that absorbed heterogeneities in the orogenic flow and maintained the network connectivity. The Eburnean example shows that mapping of finite strain and shear zone patterns from tectonic fabric trajectories is a valuable tool to infer orogen-scale kinematics and calls for caution in using local phase approach in structural analysis without considering the “bigger picture”. A new structural correlation between the West African craton and the Guiana shield provides a tectonic framework for the studied shear patterns at the scale of the Eburnean-Transamazonian orogen.

Supplementary material

Supplementary file. Structure and gold deposits of the Birimian basement of Burkina Faso.

The Supplementary Material is available at <https://www.bsgf-journal.org/10.1051/bsgf/2020001/olm>.

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